

1 Future variability of wave energy in the Gulf of Oman using a high resolution
2 CMIP6 climate model

3 Mamoud Pourali ^{a,*}, Mohamad Reza Kavianpour ^a, Bahareh Kamranzad ^{b,c,d}, Mohamad Javad Alizadeh^e

4 ^a Faculty of Civil Engineering, K.N. Toosi University of Technology

5 ^b Hakubi Center for Advanced Research, Kyoto University, Yoshida Honmachi, Sakyo-ku, 606-8501,
6 Kyoto, Japan

7 ^c Graduate School of Advanced Integrated Studies in Human Survivability (GSAIS), Kyoto University,
8 Yoshida-Nakaadachi 1, Sakyo-ku, Kyoto 606-8306, Japan

9 ^d Department of Physics, Faculty of Natural Sciences, Imperial College London, London, SW7 2AZ,
10 United Kingdom

11 ^e Caspian Sea National Research Center, Water Research Institute, Tehran, Iran

12
13 Corresponding author's email: m.pourali@email.kntu.ac.ir

14 **Abstract**

15 There is a worldwide compromise toward increasing the proportion of renewable energy in future
16 electricity production to mitigate the impacts of greenhouse gases. This study explores the
17 sustainability of wave energy resources in the northern part of the Gulf of Oman, considering the
18 impact of climate change using a Shared Socio-economic Pathway (SSP5-8.5) representing a high
19 increase in CO₂ concentration by 2100. Near-surface wind speed dataset from a high-resolution
20 CNRM (CNRM-CM6-1-HR) global climate model was employed to force a third-generation wave
21 model. A novel statistical bias-correction technique was developed based on Weibull distribution
22 to generate high-resolution input wind for the wave model, and various criteria were employed to
23 assess the sustainability of the wave energy in the study area. Comparing future projections of
24 wave energy under SSP5-8.5 with those of historical simulations demonstrated the sustainability
25 of the wave resources in the study area. The methodology of utilizing multiple criteria assessments,
26 including accessibility, availability, and exploitable storage of wave energy predicts an increase
27 ranging from 21 to 45% in the future wave power under a high emission scenario.

29 **Keywords:** Wave energy, Climate change, Weibull-based bias-correction, Gulf of Oman,
 30 Emission scenario

31

32 **Nomenclature**

33 *General:*

34	SSP	Shared Socio-economic Pathway
35	WEC	Wave Energy Converter
36	GCM	Global Circulation Model
37	IPCC	Intergovernmental Panel on Climate Change
38	CMIP6	Coupled Model Intercomparison Project Phase 6
39	SDGs	Sustainable Development Goals
40	ECMWF	the European Centre for Medium-Range Weather Forecasts
41	CNRM	Centre National de Recherches Meteorologiques
42	ERA	ECMWF Re-Analysis reanalysis
43	SW	Spectral Waves
44	<i>Downscaling:</i>	
45	$f(w)$	Weibull probability distribution on wind components
46	We+	Existing bias-correction technique based on Weibull distribution
47	We*	Proposed bias-correction technique based on Weibull distribution
48	A	Scale parameters of the Weibull distribution
49	k	Shape parameters of the Weibull distribution
50	$A'(i)$	Modified scale parameters in the month i for the future period

51	$k'(i)$	Modified shape parameters in the month i for the future period
52		
53	<i>Hydrodynamic:</i>	
54	W	Speed of wind (m/s)
55	u	Eastward wind speed (m/s)
56	v	Northward wind speed (m/s)
57	H_s	Significant wave height (m)
58	T_p	Peak wave period (s)
59	C_d	White capping
60	k_n	Bed friction, Nikuradse roughness (m)
61	N	Action density
62	σ	Frequency (hz)
63	C	Propagation velocity (m/s)
64	θ	Wave direction (rad)
65	x	Spatial coordinate in eastward direction (m)
66	y	Spatial coordinate in northward direction (m)
67	t	Time (s)
68	S	Source term
69		
70	<i>Energy:</i>	
71	P	Wave power (kW/m)
72	H_s	Significant wave height (m)

73	T_e	Energy period (s)
74	P_{ave}	Mean power (kW/m)
75	E_t	Total storage of wave energy per area (kWh/m)
76	E_e	Exploitable storage of wave energy per area (kWh/m)
77	<i>Exploitability %</i>	E_t relative to E_e
78	<i>SVI</i>	Seasonal variability index
79	<i>MVI</i>	Monthly variability index
80		
81	<i>Statistic:</i>	
82	<i>Model Skill</i>	Agreement index
83	<i>SI</i>	Scatter index
84	<i>CC</i>	Correlation coefficient
85	<i>BIAS</i>	Systematic differences between results and facts
86	<i>RMSE</i>	Root mean square error
87	X_m	Value of actual time series
88	X_p	Value of estimated time series
89	n	Number of data

90

91 **1. Introduction**

92 Wave energy as a clean and renewable energy resource is researched a lot because the majority of
 93 the human population lives in the vicinity of oceans and seas. However, it has not been
 94 commercialized yet and is still in progress. The selection of appropriate locations for energy
 95 extraction can play an important role in the efficiency of the Wave Energy Converters (WECs)

96 [1]. Formerly, it was deemed that locations with higher mean annual wave power are desirable,
97 while less attention was paid to temporal variation of the power. Recently, it was found that
98 locations with lower energy, but higher temporal stability are more favorable in terms of wave
99 energy extraction compared to those with higher energy but less temporal stability [2]. [3] and [4]
100 explored the performance of different wave energy converters at different installation depths and
101 at different locations concluding that not only the suitable WEC type, but also the optimal design
102 varies for different locations and depths. Along with intra-annual fluctuations, the future variation
103 of wave energy resources due to changing climate may significantly affect the viability of the
104 power extraction due to the increase in the greenhouse gas concentration. Therefore, many studies
105 have been devoted to investigating the impacts of climate change on renewable energy resources
106 [5–9].

107 Considering climate change impacts on renewable energy resources, it is important to take into
108 account the trends in the current/past wave power and explore the possible future variations.
109 Assessing the long-term behavior of wave conditions, [10] reported an increase in global wave
110 power by 0.4% per year from 1984 to 2017. The results align with the other studies that
111 demonstrated increasing wave energy trends for different regions [4] [11]. Global Circulation
112 Models (GCMs) provide the future projections of various atmospheric and oceanic variables.
113 These models are developed and run globally, considering several possible future scenarios.
114 Following the latest updates of the Intergovernmental Panel on Climate Change (IPCC) released
115 as Coupled Model Intercomparison Project phase 6 (CMIP6), the model projecting climatic
116 variables is developed according to Shared Socio-economic Pathways (SSPs). Wind resource
117 evolution under different CMIP6 climate change scenarios in Europe and North America has been
118 investigated by [12] and [13]. However, wave characteristics are not projected in GCMs, directly.
119 Thus, the wind field obtained from GCMs is utilized as a driving force to wave models to generate
120 the wave characteristics and subsequently, investigate the wave resources. Since GCMs commonly
121 run at a coarse spatial resolution which may not cover the local fluctuations, bias-correction or
122 downscaling processes may be required for regional studies. Regionalization of GCM simulations
123 is a common practice to achieve more reliable results consistent with the conditions governing the
124 area.

125 Regression models, artificial intelligence models, quantile mapping models and Weibull
126 distribution models are all already used for wind field modifications [14], [15]. Based on a
127 comparison of the performance of different statistical downscaling techniques, including
128 multiplicative shifting, quantile mapping, support vector regression, and Weibull-based
129 techniques, the Weibull-based technique outperformed them all [15]. This method is based on the
130 probability distribution of the wind components and simultaneously modifies wind speed and
131 direction. Moreover, the method does not disturb the sequential order of time series, which are
132 important in wave power projections. However, there is still a need for improvement and progress
133 to enhance the efficiency and accuracy of the method. In fact, since the Weibull equation is
134 inherently in a multiplicative form because of a coefficient called the shape factor, using the
135 multiplicative relation will improve the estimates compared to the additive method. Because of
136 their strong dependence and sensitivity on the wind field, choosing and applying a suitable bias-
137 correction technique is crucial to the accuracy and viability of wave projections. [16].

138 In order to select the appropriate locations and WECs for wave energy extraction, it is vital to
139 consider various criteria in line with the sustainability of the resources. The suitability of less
140 powerful but more consistent wave resources has been addressed by [17] and [18]. In addition, the
141 authors discussed different factors influencing the optimal site selection for wave energy farms.
142 Assessment of renewable energy resources in Iran with a focus on marine resources has revealed
143 many energy hotspots with a high potential for marine energy development [19]. The Gulf of
144 Oman, with its vicinity to the Indian Ocean and the swells traveling from the Southern Ocean, has
145 the potential to supply part of the energy demands for the population living in the coastal areas.
146 More importantly, there are several coastal villages far from the cities where access to electricity
147 is limited. In addition, with its proximity to the Indian Ocean and access to open waters, the Gulf
148 of Oman is an attractive location for various industrial projects and developments, whereas the
149 complex sea state is highly affected by both Shamal and monsoon winds from the west and south,
150 respectively [16] and makes the spatio-temporal analysis required. Considering the local energy
151 resources such as wave energy, it can be considered a cost-effective and sustainable choice for the
152 electricity supply. There have been previous studies on the investigation of wave energy resources
153 in the Gulf of Oman, such as [20,21], in which the analysis has been done based only on the mean
154 wave energy values and their variability on a monthly scale. In addition, [22] proposed a multi-

155 criteria approach to select the most appropriate combination of wave energy convertor and location
156 in several stations in the Caspian Sea, the Persian Gulf, and the Gulf of Oman. They used different
157 factors, including exploitable storage of energy, accessibility, availability, energy production,
158 design wave height, and intra-annual variation of the resources. However, their analysis has been
159 done only based on the historical simulation of wave characteristics and lacks long-term changes
160 due to climate change.

161 The main objective of this study is to develop an accurately modeled wave field to estimate the
162 future wave power conditions in the northern part of the Gulf of Oman. For this purpose, the wind
163 dataset of SSP5-8.5 derived from a CMIP6 model is modified by developing a Weibull distribution
164 based. Afterward, the modified wind field is used to run the third-generation model MIKE 21
165 Spectral Waves (SW) [23] to obtain wave characteristics. Following the proposed approach based
166 on different criteria, the sustainability of wave energy resources under the impact of climate change
167 is investigated in four locations. In section 2, the methodology includes study area, calibration,
168 climate data, numerical wave model characteristics, the proposed downscaling technique,
169 modeling procedures, and wave power computations. Results are presented in section 3, while a
170 [discussion is given in section 4](#). The main findings and conclusions are summarized in section 5.

171

172 2. Materials and methods

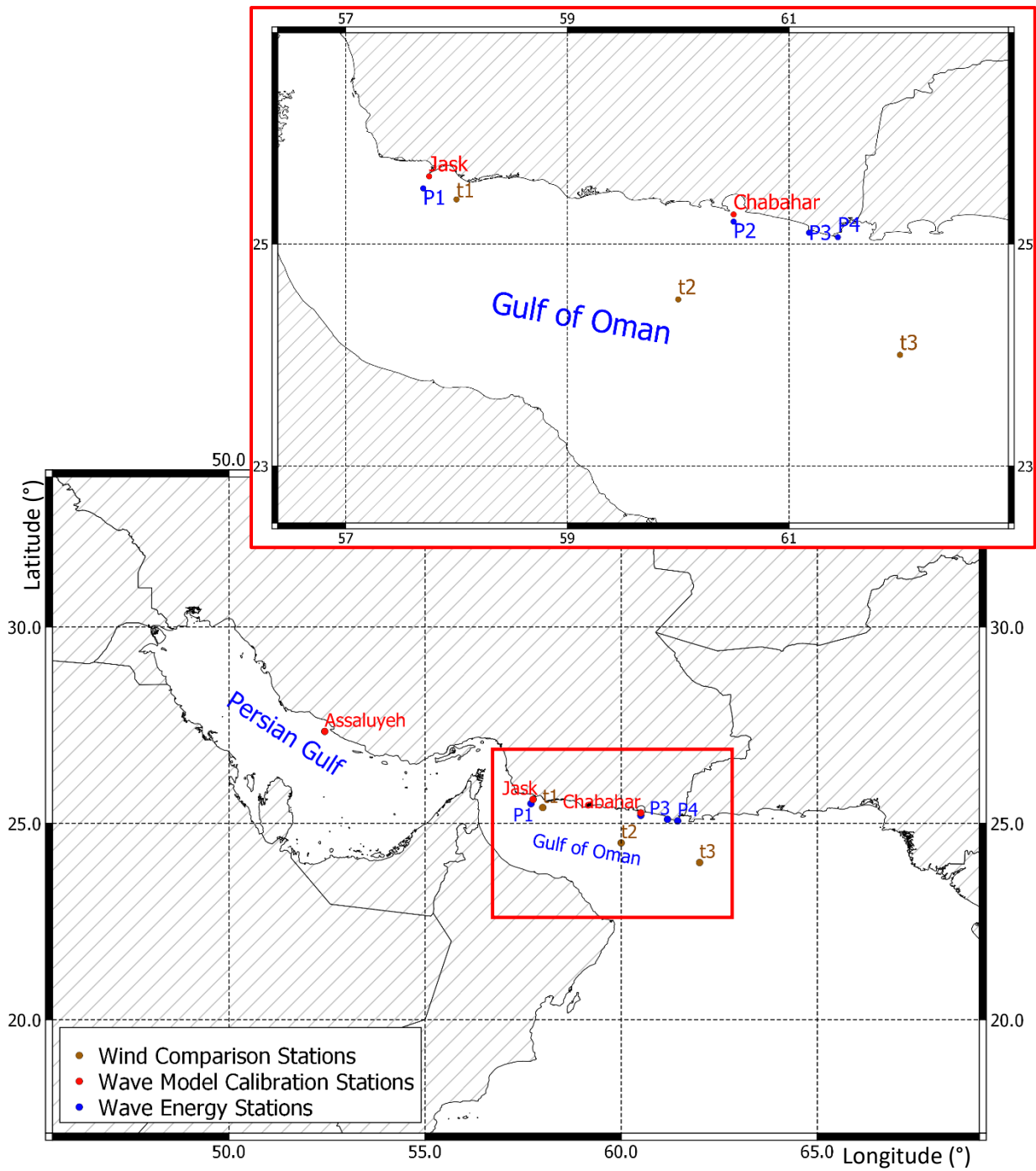
173 The methodology of this study is described as: selection of specific locations in the study area,
174 collecting the dataset, modification of wind field derived from the climate model, performance
175 evaluation of the numerical wave model, calculation of wave power, and assessing the climate
176 change impact on wave energy and temporal stability analysis. In climate change impact studies,
177 it is common to investigate future projections in either the near or far future (e.g. [24]). Considering
178 the goal of this study to focus on the sustainability of resources in the long-term and in order to
179 further compare the results with additional relevant studies, a period of 100-year has been selected.
180 Four points in the northern Gulf of Oman have been chosen for site selection. Two datasets,
181 including the near-surface wind speed of ERA5 reanalysis dataset [25] from 1981 to 2000, and
182 also historical wind speed (1981-2000) and future (2081-2100) scenario of SSP5-8.5 from CMIP6
183 CNRM climate model [26], were used. The dataset was selected since it has the highest spatial

184 resolution among the available GCMs. The area is affected by various climates, such as monsoons
185 from the Indian Ocean and shamal wind from the northwest to some extent. Hence, high-resolution
186 wind data plays an important role in not only generating higher accuracy wave height and period
187 but also a correct propagation direction. The CNRM data are bias-corrected (considering ERA5 as
188 the reference dataset for the historical period) by developing an improved statistical technique
189 based on Weibull distribution. The modified data of wind components are used to run the
190 numerical wave model for both historical and future periods. To estimate the wave energy, the
191 numerical wave model (MIKE21 SW) results are used. Finally, different analyses are carried out
192 to explore climate change impacts and wave energy variability in the long term at the selected
193 locations.

194

195 2.1. Study area and data

196 The selected locations in this study include four spots in the northern Gulf of Oman in the
197 nearshore areas of Iranian waters. The Chabahar free trade-industrial zone is located there where
198 it is experiencing fast growth and subsequently increasing demand for energy production.
199 Moreover, there are rural and urban regions nearby the Gulf, where the development of local
200 energy sources can be economically viable and reduce the cost of long-distance energy transfer
201 from a power plant to the consumption units. Regarding the population and the importance of these
202 regions, four points (P1 to P4) as candidate locations are selected for wave energy analyses, as
203 illustrated in Fig. 1 and Table 1. Moreover, distance from the residential areas to the energy sources
204 was also considered in selecting the locations. It should be noted that the computational domain in
205 both wind calibrating and wave simulation includes a rectangular area covering the longitudes
206 47.5°E-74.0°E and latitudes 15.0°N-30.5°N. It covers the Gulf of Oman above 15° latitude to the
207 Persian Gulf. The target area of the modeling is shown in Fig. 1.



208

209

Fig. 1. Study area, wind comparison spots, and selected wave energy locations

210

211

212

Table 1. Locations of the selected spots for wave analysis

Point	Location	Long. (°)	Lat. (°)	Depth [m]
P1	Jask	57.7	25.5	100
P2	Chabahar	60.5	25.2	55
P3	Beris	61.18	25.1	14
P4	Pasabandar	61.44	25.06	30

213

214

2.2. Reanalysis and climate model datasets

215

216

217

218

219

220

221

222

223

224

225

226

227

228

229

230

231

232

233

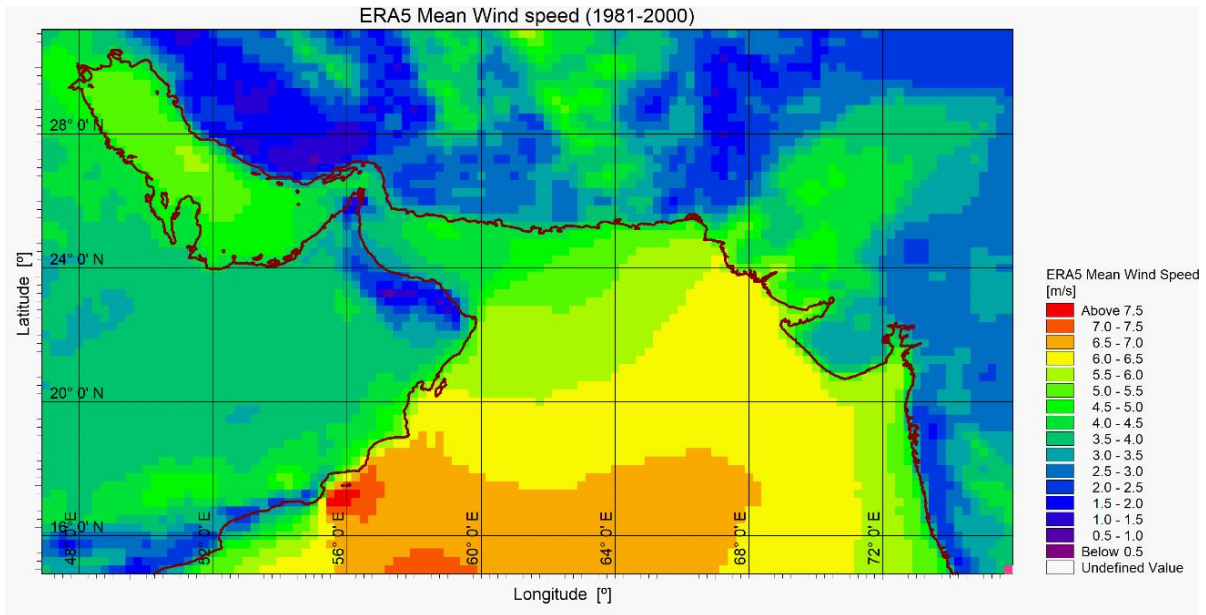
234

Two sets of data have been used for the wave climate investigation in the study area. Eastward and northward wind components of ERA5 reanalysis with a spatial resolution of $0.25^\circ \times 0.25^\circ$ and temporal resolution of 1 hour have been downloaded from <https://cds.climate.copernicus.eu/> and were used as the reference data for bias-correction of the wind field obtained from the CMIP6 GCM. The data was obtained for the whole computational domain for 20 years, from 1981 to 2000. Skill assessment and validation of ERA5 reanalysis dataset against measurements and altimeter observations indicated its efficiency for wind field simulation over the study area [27][28]. Moreover, several studies in different areas applied ERA5 and its previous versions developed by ECMWF, as reference data for downscaling/bias-correction of GCMs on a local scale [21,29–31].

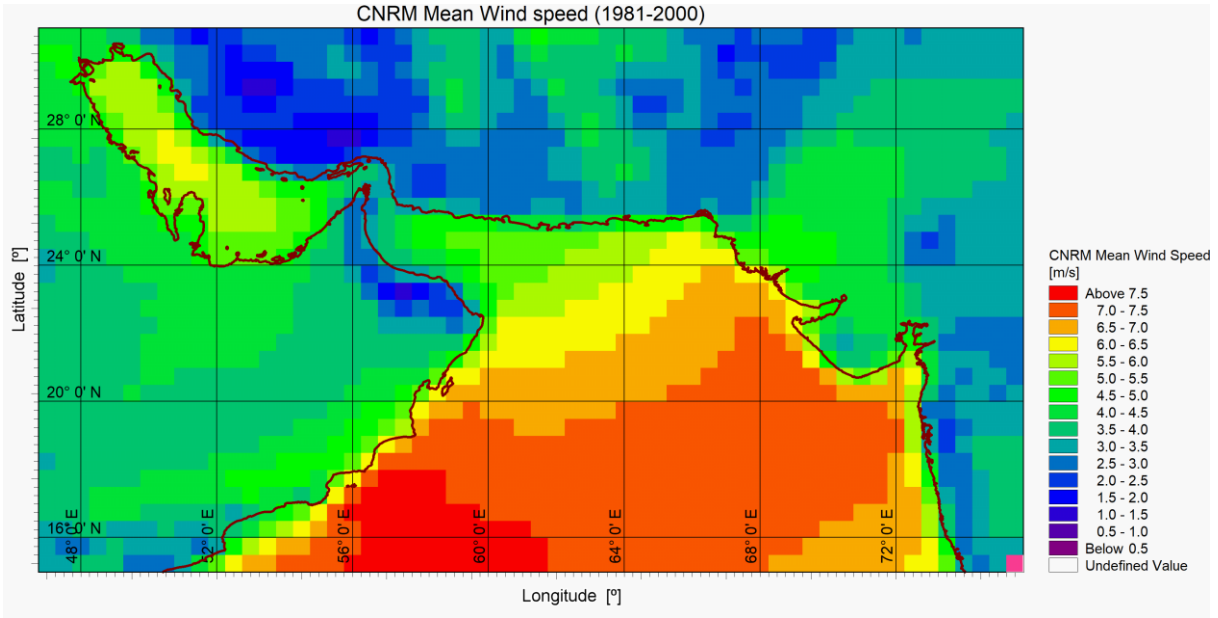
Wind speed characteristics as an input to the wave model are necessary to evaluate future variability in wave power. In CMIP6 models, the climate variables for the future period are available for different SSPs. In this regard, four SSPs; SSP1-2.6, SSP2-4.5, SSP3-7.0, and SSP5-8.5 are used for climate change studies [32]. According to the purpose of this study which is to explore the future variability of the wave power under climate change impacts, the wind speed simulation of the scenario SSP5-8.5 is considered for future wave power projection. SSP5-8.5 is a high emission scenario which is an updated version to CMIP5 RCP8.5, indicating radiative forcing of 8.5 W/m^2 by 2100, and it assumes a fossil-based economy in the future. SSP5-8.5 is a pessimistic socio-economic pathway scenario exacerbating future changes due to a rapid increase in CO_2 concentration. Finally, the worst-case scenario (SSP5-8.5) is expected to result in the highest variability in the future atmospheric conditions than the other scenarios. Thus, it is

235 important to consider such variations for future planning and management. In addition, many
 236 studies on climate change impacts on wave climate have used RCP8.5 (the equivalent of SSP8.5
 237 in CMIP5 models) [33], [24], and [8]. Hence, this study focuses on using this scenario for future
 238 projections.

239 The next step is to select an appropriate GCM with reliable performance in the study area,
 240 considering its spatio-temporal resolution and coverage. To that end, GCM evaluation is a common
 241 task to find the climate model with the best performance for a particular area [34]. Following the
 242 previous studies and also due to the high spatial resolution, the CNRM-CM6-1-HR outputs have
 243 been selected in this study [35]. Thus, near-surface wind speed of the historical (1981-2000) and
 244 future period (2081-2100) from the CNRM model have been obtained and assessed. The data has
 245 a temporal resolution of 3 hours and spatial resolution of $0.5^{\circ} \times 0.4993^{\circ}$ for both historical and
 246 future periods. Fig. 2 illustrates the mean annual wind speed for ERA5 and CNRM-CM6-1-HR
 247 model over the computational domain from 1981 to 2000. This figure shows a bias in the wind
 248 speed simulations obtained from the GCM compared with ERA5, especially in the southern parts
 249 of the Gulf of Oman when it is connected to the Indian Ocean.



250 a)



251 b)

252 Fig. 2. Mean annual wind speed for historical period (1981-2000) a) ERA5, b) CNRM

253 As shown in Fig. 2, the GCM used in this study (CNRM) needs to be modified to resemble the
 254 wind field on the regional scale properly, and modifications are required in order to increase the
 255 accuracy of the wind data. Table 1 presents the statistical characteristics of the ERA5 and CNRM
 256 wind speed simulations in the selected locations (P1 to P4) for the historical period. The statistics
 257 in Table 2 show that the CNRM model overestimates the wind speed for the selected locations in
 258 terms of minimum, maximum, average, and standard deviation calculated for 20 years. The
 259 standard deviation is also higher for the wind speed derived from CNRM. Hence, bias correction
 260 is required before applying the CNRM outputs for wave climate projection, assuming that the same
 261 correction method is applied to the future dataset.

262

263 Table 2. Statistics of the wind speed data from ERA5 and CNRM (1981 to 2000)

Station	ERA5				CNRM			
	Min.	Max.	Avg.	Std.	Min.	Max.	Avg.	Std.
Jask	0.06	14.852	4.645	2.236	0.139	18.169	4.946	2.634
Chabahar	0.068	15.618	4.144	1.904	0.087	24.059	4.471	2.302
Beris	0.146	15.087	4.053	1.855	0.149	23.824	4.314	2.221

Pasabandar	0.117	14.515	4.070	1.881	0.183	24.070	4.405	2.267
------------	-------	--------	-------	-------	-------	--------	-------	-------

264

265 2.3. The proposed bias-correction technique

266 In this study, an improved version of the Weibull-based technique as a statistical bias-correction
 267 approach is employed to modify the wind field. The initial version of this technique was proposed
 268 by [15] to fit Weibull probability distribution on wind components through a data-transformation
 269 process. Generally, Weibull distribution can be defined as:

270
$$f(W) = \frac{k}{A} \left(\frac{W}{A}\right)^{k-1} \exp \left[-\left(\frac{W}{A}\right)^k\right]$$

271 (1)

272 Where W is the speed of wind and A and k are called scale and shape parameters of the distribution.
 273 Following Eq. (1), it can be found that the distribution is dependent on these two parameters and
 274 having the time series of wind speed, the main practice is to find them through the standard
 275 deviation method or alternatively by means of the maximum likelihood method [36]. Weibull
 276 (shape and scale) parameters of the historical datasets of ERA5 and CNRM model for each
 277 gridpoint are compared, and their ratio is used as a correction factor to modify the wind field in
 278 the future period for SSP5-8.5. These coefficients can be obtained as Eqs. (2-3) and employed for
 279 future wind field modifications as Eqs. (4-5).

280
$$D_{A(i)} = A_{ERA5(i)}^{His} / A_{CNRM(i)}^{His} \tag{2}$$

281
$$D_{k(i)} = k_{ERA5(i)}^{His} / k_{CNRM(i)}^{His} \tag{3}$$

282
$$A'_{CNRM(i)}^{Future} = A_{CNRM(i)}^{Future} * D_{A(i)} \tag{4}$$

283
$$k'_{CNRM(i)}^{Future} = k_{CNRM(i)}^{Future} * D_{k(i)} \tag{5}$$

284 Where $A(i)$ and $k(i)$ are the modified scale and shape parameters in the month i for the future
 285 period. In comparison with the previously proposed method [15], in which the difference was used
 286 as the correction coefficient, this study uses the proportion of the Weibull parameters in the
 287 historical period for the two datasets. An attempt is made to examine and compare the
 288 multiplicative form against the difference form as an alternative.

289 The efficiency of the proposed bias-correction technique depends on the assumption that the wind
290 field in the region follows a Weibull distribution. As previous studies showed, in the study area,
291 Weibull distribution fits suitably with the wind data. Therefore, a distribution-based model can be
292 beneficial to both modifying the data statistics and the probability distribution of the data.
293 Moreover, the application of the Weibull distribution is limited to positive values of $W \geq 0$ where
294 wind components (u and v representing eastward and northward wind speeds) can be either
295 positive or negative depending on the direction they show. Therefore, the data transformation
296 process should be implemented on the wind components before fitting the distribution. The scheme
297 for data transformation is adding the absolute of the strongest wind components (u, v) with a
298 negative sign to the time series of the same point and the same component. Although the
299 transformation may introduce an irreversible error, the efficiency of this method is still
300 advantageous to the traditional approach of applying the Weibull distribution on the wind speed
301 (W) directly. This is mainly due to wind speed and direction modification when wind components
302 are improved. It is noted that the suitability of the Weibull distribution on the transformed wind
303 components is checked throughout the bias-correction process. More information concerning the
304 validity of the Weibull-based model can be found in [15].

305 The Weibull distribution is fitted on wind components (u, v) individually. The shape and scale
306 parameters are computed for the historical period of the reference and GCM wind components.
307 Afterward, the correction factor is obtained from the historical datasets (Eqs. 2 and 3) and
308 multiplied by the corresponding shape and scale factors values in the future periods (Eqs. 4 and
309 5). Finally, wind components are modified and subsequently de-transformed to their original
310 ranges utilizing the modified parameters and inverse Weibull distribution. This procedure is
311 repeated for all the gridpoints, one by one, using a distributed scheme to modify the wind field
312 over the whole computational domain. The modified wind field can be used as input for the
313 numerical wave model to generate the wave characteristics and, consequently, the wave power.

314

315 2.4. Numerical wave model

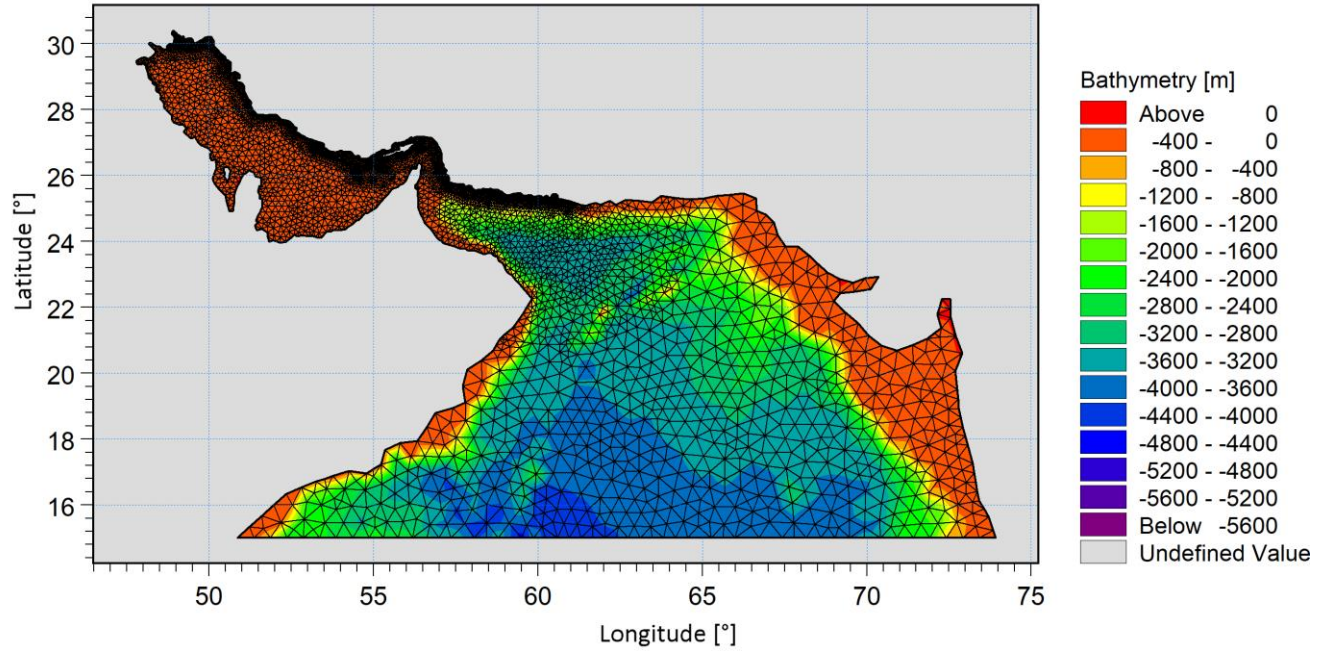
316 In this study, MIKE 21 SW, a third-generation spectral numerical model based on unstructured
317 meshes, is employed to simulate wave characteristics in historical and future periods. The model

318 has been successfully applied for nearshore and offshore wave modeling [37–40]. It uses a spectral
 319 action balance equation considering the growth, decay, and transformation of wind-generated
 320 waves and swells in offshore and coastal areas to simulate wave evolution [23]. The basic equation
 321 can be formulated either in the Cartesian coordinates for small-scale applications and polar
 322 spherical coordinates for large-scale applications. Thus, for this study, the wave action balance
 323 equation is written in spherical coordinates as:

$$324 \quad \frac{\partial N}{\partial t} + \frac{\partial}{\partial \phi} C_{\phi} N + \frac{\partial}{\partial \lambda} C_{\lambda} N + \frac{\partial}{\partial \sigma} C_{\sigma} N + \frac{\partial}{\partial \theta} C_{\theta} N = \frac{S}{\sigma} \quad (6)$$

325 Where N is the action density, t is time, ϕ is the latitude, λ is the longitude, σ is frequency, C is the
 326 propagation velocity, and θ is wave direction in a nautical convention. The first three terms on the
 327 left side of Eq. (6) show temporal and spatial variation of N . The fourth term represents how the
 328 relative frequency changes as depth changes, while the fifth term represents refraction caused by
 329 depth and currents. The total source and sink function is represented by S reflecting the effects of
 330 the generation by wind, dissipation (by white-capping, depth-induced wave breaking, and bottom
 331 friction) and nonlinear wave-wave interactions [41]. Since MIKE 21 SW is embedded in two
 332 different formulations, including fully spectral and directional decoupled parametric formations,
 333 in this study, the fully spectral formulation is used based on the wave action conservation equation
 334 [23,42]. Figure 3 shows the computational wave domain, the size of the elements, and the
 335 bathymetry of the prepared wave model.

336



337

338

Fig. 3. Wave modeling domain, element size, and bathymetry of the wave model

339

340 The following statistical relationships were used to evaluate the wave model. They are among the
 341 standard statistical metrics to evaluate the accuracy of the model simulations. Generally, models
 342 with relatively higher *Model Skill* and *CC* (closer to 1), but with lower *SI*, *RMSE*, and *BIAS* are
 343 desirable. Table 3 shows the results of the wave model evaluation based on these formulas in
 344 Assaluyeh, Jask, and Chabahar stations.

$$345 \quad Model\ Skill = 1 - \frac{\sum(X_p - X_m)^2}{\sum(|X_p - \bar{X}_m| + |X_m - \bar{X}_m|)^2} \quad (7)$$

$$346 \quad SI = \frac{\sqrt{\frac{1}{n} \sum(X_p - X_m)^2}}{\bar{X}_m} \quad (8)$$

$$347 \quad CC = \frac{\sum(X_p - \bar{X}_p)(X_m - \bar{X}_m)}{\sqrt{\sum(X_p - \bar{X}_p)^2 \sum(X_m - \bar{X}_m)^2}} \quad (9)$$

$$348 \quad BIAS = \sum \frac{1}{n} (X_p - X_m) \quad (10)$$

$$349 \quad RMSE = \sqrt{\frac{1}{n} \sum (X_p - X_m)^2} \quad (11)$$

350 where X_m and X_p are the actual and the estimated time series, respectively. $\overline{X_m}$ and $\overline{X_p}$ are the
 351 average of the real and simulated values. Moreover, n denotes the number of the data, and the
 352 overbar indicates the sample mean.

353

354 2.5. Factor to assess the stability and suitability of resources

355 The outputs of the numerical wave models are available as significant wave height and wave
 356 period. Hence, to calculate the wave power (kW/m), deep water approximation was utilized
 357 [43,44]:

$$358 \quad P \approx 0.49 H_s^2 T_e \quad (12)$$

359 where H_s is the significant wave height and T_e is the energy period, which is a function of spectral
 360 moments of order 0 (m_0) and -1 (m_{-1}) as follow:

$$361 \quad T_e = \frac{m_{-1}}{m_0} \quad (13)$$

362 Using the wave power value for each time step, the mean power (P_{ave}) is calculated. Also, having
 363 the total hours per year (t), the total and exploitable storages of wave energy per unit area (E_t and
 364 E_e , respectively) can be obtained as:

$$365 \quad E_t = P_{ave} \times t \quad (14)$$

$$366 \quad E_e = P_{ave} \times t_e \quad (15)$$

367 Where t and t_e represent the total hours all year round and the total hours when the energy is
 368 greater than a threshold, respectively. The annual electric power production of a WEC device is
 369 calculated via a coupling of the device's power matrix and the wave resource matrix [45][46]. In
 370 this study, the threshold is considered 2 kW/m as suggested and applied by [45,47,48]. Moreover,
 371 extreme events are excluded in the calculations considering an upper limit for significant wave
 372 height as $H_s \geq 4m$ [49]. Ratio of E_e to E_t is used as *exploitability*:

$$373 \quad \textit{Exploitability} (\%) = \frac{E_e}{E_t} \quad (16)$$

374 Since this study aims to evaluate the viability and sustainability of wave power under a high
 375 emission scenario (SSP5-8.5), intra-annual variability (monthly and seasonal variability) and
 376 availability must be considered. These factors depend on wave conditions that may change under
 377 the climate change impacts. It is worth mentioning that accessibility is another sea state-dependent
 378 factor that has been missed in this study since it was shown by [22] that accessibility reaches about
 379 100% for Govatr located in the study area. Availability represents the percentage of time in which
 380 the wave characteristics lie in the desirable range of wave energy devices to operate efficiently.
 381 The range is restricted by cut-in and cut-off values of significant wave height, which are considered
 382 0.5 m and 4 m, respectively, according to [50].

383 Seasonal and Monthly Variability Indices (*SVI* and *MVI*, respectively) are used to assess the
 384 temporal stability of both historical and future periods.

$$385 \quad \textit{SVI} = \frac{P_{S1} - P_{S4}}{P_{ave}} \quad (17)$$

$$386 \quad \textit{MVI} = \frac{P_{M1} - P_{M12}}{P_{ave}} \quad (18)$$

387 where P_{S1} and P_{M1} represent the mean wave power for the highest energy season and month,
 388 respectively, while P_{S4} and P_{M12} represent the mean wave power for the least energetic season and
 389 month, respectively. The annual mean wave power is denoted by P_{ave} . Generally, higher values of
 390 mean wave power and also lower values of *MVI* and *SVI* demonstrate more temporally stable
 391 conditions. Comparing the mentioned variables in the historical simulations with those of the
 392 future projection can provide the required information for decision-makers and authorities to have
 393 a suitable development plan considering the Sustainable Development Goals (SDGs).

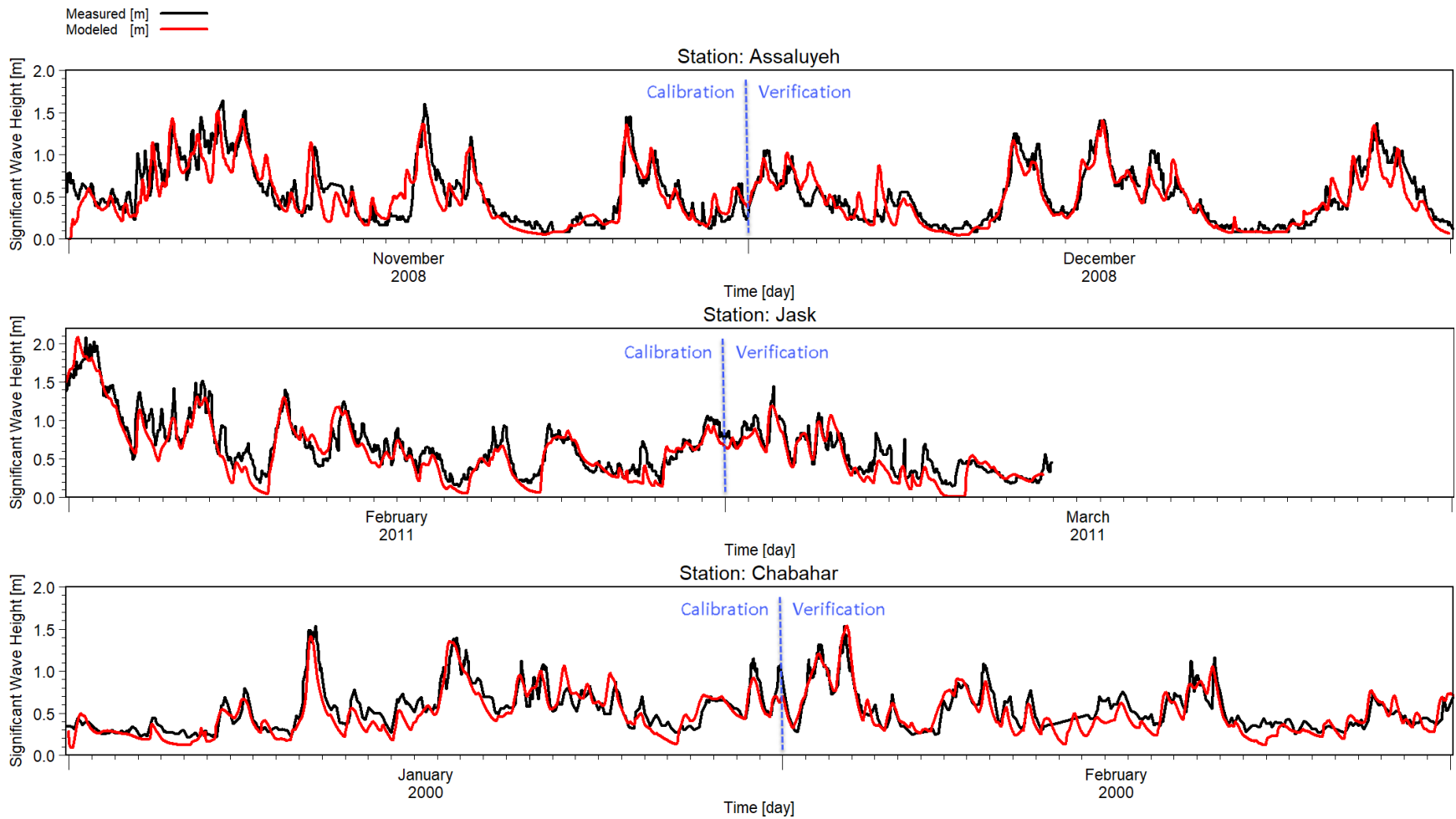
394

395 3. Results

396 3.1. Wave model calibration and validation

397 To calibrate the numerical wave model, firstly, sensitivity analyses were conducted to find the
 398 most effective parameters on the model outputs. Subsequent to the mesh size optimization through

399 the sensitivity analysis, white capping and bed friction were considered as the calibration factors.
400 In case of white capping, the two parameters of C_d and δ should be tuned accordingly. The
401 bed friction can be specified with Nikuradse roughness parameter (k_n). After trial and error, the
402 values of 2, 0.5, and 0.001 were selected for C_d , δ and Nikuradse roughness, respectively. In
403 terms of directional discretization and time formulation, 16 directions (22.5° resolution) and an
404 instationary formulation were used. The evaluation was based on the model's capability to simulate
405 the significant wave height (H_s) and peak wave period (T_p) in Assaluyeh, Jask and Chabahar where
406 the observational records are available, while the wind field of ERA5 as the input to force MIKE
407 21 SW. Fig. 4 illustrates the time series of the numerical wave model versus the observational data
408 for the three stations.



409

410

411

412

413

Fig. 4. Time series of the significant wave height (H_s) for the numerical model and observations in Assaluyeh, Jask, and Chabahar stations

414

415 Table 3. Simulation statistics of the significant wave height (H_s) in meters and peak wave period
 416 (T_p) in seconds

Buoy	Assaluyeh		Jask		Chabahar	
Data Period	Nov & Dec-2008		Feb & Mar-2011		Jan & Feb-2000	
Parameter	T_p	H_s	T_p	H_s	T_p	H_s
<i>Model Skill</i>	0.96	0.93	0.86	0.90	0.94	0.95
<i>SI</i>	0.19	0.35	0.17	0.34	0.11	0.22
<i>CC</i>	0.92	0.86	0.75	0.84	0.90	0.91
<i>BIAS</i>	-0.14	-0.03	-0.03	-0.08	0.02	-0.05
<i>RMSE</i>	0.76	0.18	0.76	0.21	0.13	0.15

417

418 According to Fig. 4 and Table 3, the observed and simulated wave heights coincide at all three
 419 stations. Values of error indices confirm the efficiency of the model. Therefore, the calibrated
 420 model can be employed to simulate the significant wave height in the area for a longer term.

421

422 3.2. Bias-correction of the wind field

423 The performance of the numerical wave models strongly depends on the accuracy of the wind field
 424 and its consistency with the local conditions. As mentioned earlier, bias correction is required for
 425 modifying the wind field obtained from the climate model. In this section, the performance of the
 426 proposed bias-correction technique has been provided in Table 4. Moreover, the results are
 427 compared with the recently developed bias-correction technique but with newly proposed
 428 modification criteria, as explained in section 2.3. It is pointed out that $We+$ stands for the existing
 429 bias-correction technique based on Weibull distribution, and We^* is the proposed method in this
 430 study which is a modified version of the $We+$ technique. In this comparison, sample points of t1,
 431 t2, and t3 belong to the area close to the Strait of Hormuz in the Gulf of Oman (Fig. 1).

432

433 Table 4. Performance of the developed bias-correction technique for the wind speed (compared
434 with ERA5)

Location	t1			t2			t3		
	Long.(°)		Lat. (°)	Long. (°)		Lat. (°)	Long. (°)		Lat. (°)
	58		25.4	60		24.5	62		24
Technique	Raw	We+	We*	Raw	We+	We*	Raw	We+	We*
<i>SI</i>	0.80	0.667	0.666	0.78	0.629	0.629	1.24	0.589	0.589
<i>CC</i>	0.06	0.195	0.196	0.03	0.155	0.156	0.016	0.463	0.467
<i>BIAS (m/s)</i>	0.58	-0.169	-0.169	0.66	0.043	0.040	2.22	-0.108	-0.113
<i>RMSE (m/s)</i>	2.79	2.313	2.311	3.39	2.716	2.714	3.96	1.881	1.872

435
436 Along with the metrics presented in Table 4, the capability of the bias-correction technique to
437 suitably resemble extreme events is of great importance and useful for many practical applications.
438 In this regard, 95th and 99th percentile as indicators for extreme values of the wind speed for a 20-
439 year period were calculated for the reference data and modified wind field obtained by
440 implementing the bias-correction technique on the data from the climate model. Table 5 compares
441 the efficiency of the proposed bias-correction technique with the existing technique in terms of
442 95th and 99th percentiles. It is pointed out that ERA5 is considered as the reference dataset to
443 evaluate the performance of the bias-correction techniques.

444
445 Table 5. Performance of the bias-correction techniques to estimate extreme wind speeds

95 th and 99 th percentile of wind speed								Absolute difference with ERA5			
t1								t1			
95 th percentile				99 th percentile				95 th percentile		99 th percentile	
ERA5	Raw	We+	We*	ERA5	Raw	We+	We*	We+	We*	We+	We*
6.699	8.2	6.771	6.751	8.025	10.67	8.584	8.566	0.0725	0.0524	0.5589	0.5411
t2								t2			
95 th percentile				99 th percentile				95 th percentile		99 th percentile	
ERA5	Raw	We+	We*	ERA5	Raw	We+	We*	We+	We*	We+	We*

7.838	9.90	8.319	8.307	9.956	12.8	10.538	10.555	0.4819	0.4697	0.5814	0.5990
t3				t3				t3			
95th percentile				99th percentile				95th percentile		99th percentile	
ERA5	Raw	We+	We*	ERA5	Raw	We+	We*	We+	We*	We+	We*
6.381	10.4	6.552	6.533	8.016	12.9	8.584	8.513	0.1711	0.1523	0.5674	0.4967

446

447 As observed from Table 5, the GCM simulations (here indicated as Raw) overestimate wind speed
 448 in terms of 95th and 99th percentiles for all the selected locations. Comparing the extreme values
 449 calculated from the modified wind field by the two bias-correction techniques indicates that the
 450 proposed method slightly outperforms its former counterpart. Moreover, there is consistency
 451 between the extreme values obtained from the reference data with those of the corresponding
 452 values calculated from the modified wind field of the climate model. Considering the absolute
 453 difference between the extreme values of ERA5 and the modified wind fields using the bias-
 454 correction technique, slightly lower errors (improved accuracy) are represented by the proposed
 455 bias-correction technique than the previous method. The proposed bias-correction technique
 456 remarkably improves climate models' wind speed in which the modified wind field is compatible
 457 with the reference data. As stated by [29], the additive correction of the Weibull parameters leads
 458 to significant improvement of the wind direction. To assure suitability of the developed method in
 459 this study, which is the multiplicative correction of the parameters, wind roses for the three
 460 locations are depicted for reference data and modified wind field from the climate model by
 461 additive and multiplicative correction factor (Fig. 5).

462

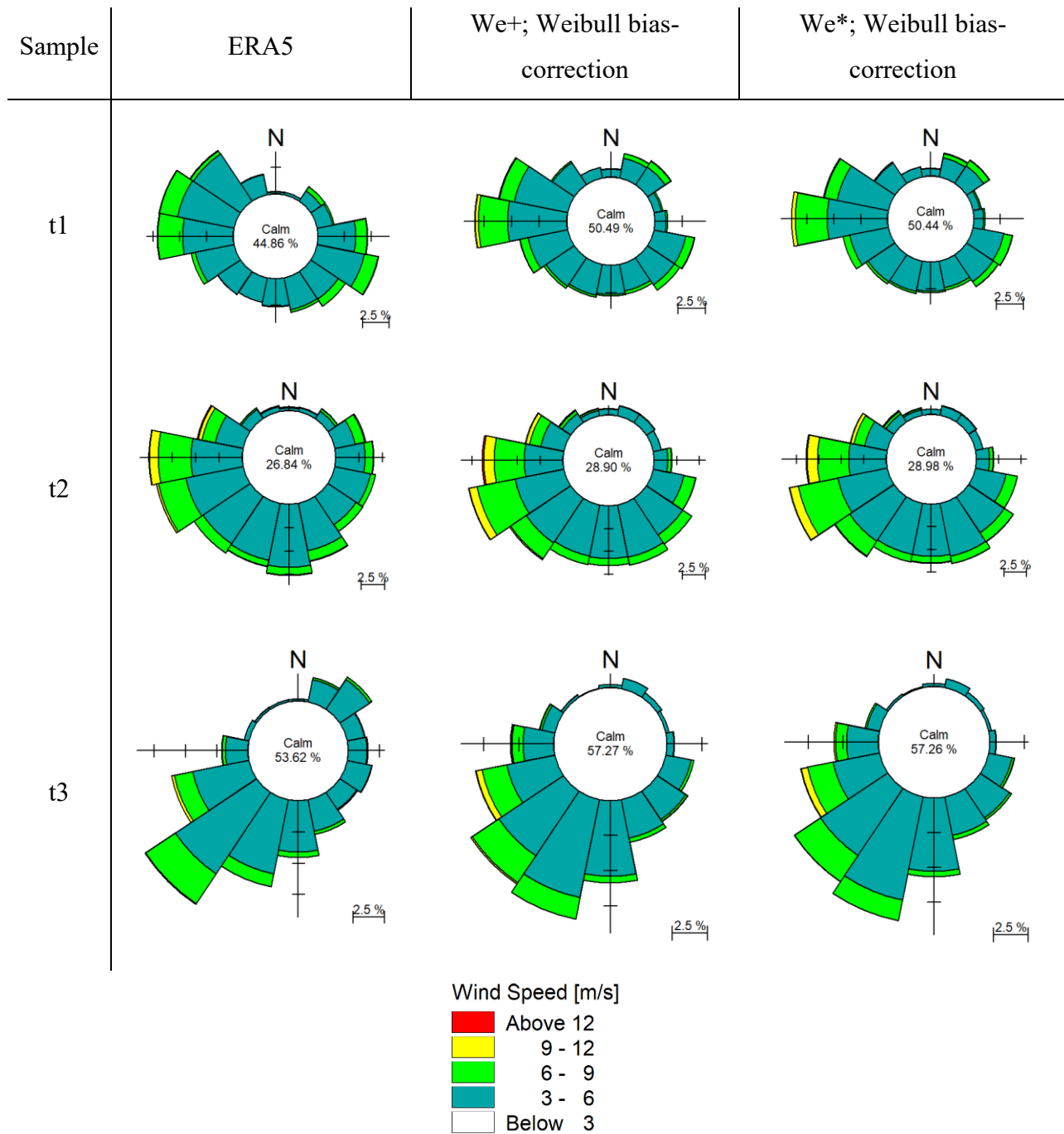
463

464

465

466

467



468

Fig. 5. Wind roses for different locations

469

470 Following Fig. 5, it is inferred that the directional distribution of the wind speed obtained from the
471 two bias-correction techniques is nearly similar. Furthermore, similarity in magnitude and
472 directions of wind speed for the reference data and modified data reveal the efficiency of the bias-
473 correction technique. Considering the capability of the bias-correction technique to modify
474 directional and magnitude distributions of the wind for the historical period as a validation stage,
475 it can also be applied to the future wind field. Afterward, they can be employed for wave power
476 projection to assess climate change impacts.

477

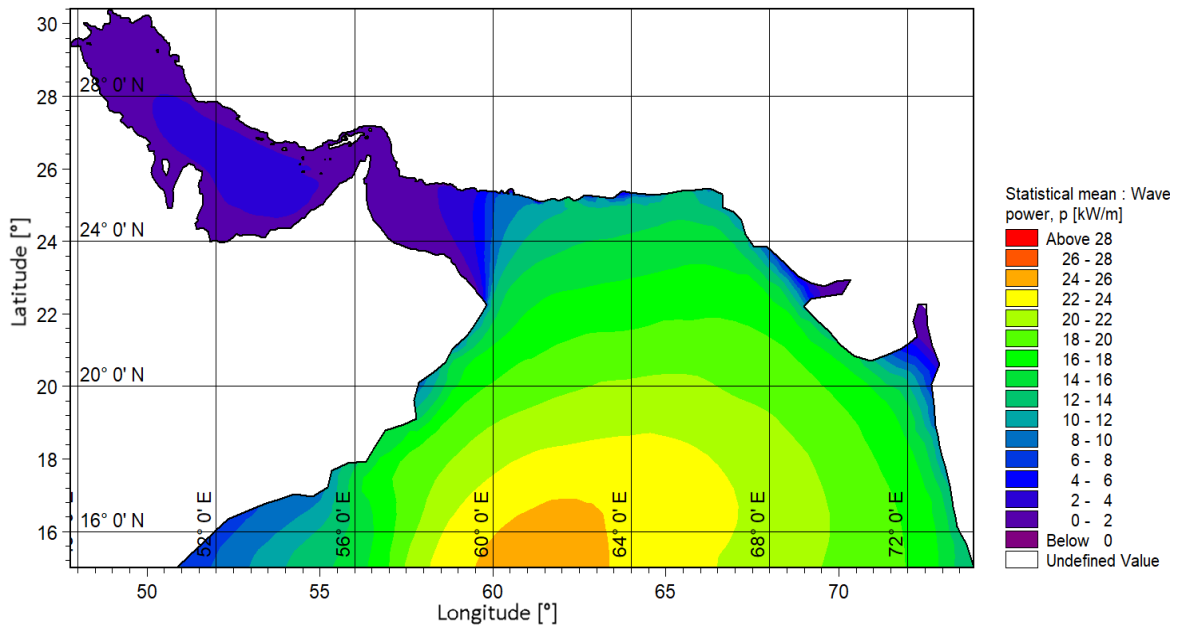
478 3.3. Historical and future exploitable wave energy

479 The calibrated wave model is employed for the future wave power projection. In this regard, the
480 correction coefficient obtained for the historical wind field is applied to modify the future wind
481 field representing the high emission SSP5-8.5 of the CNRM model. Climate change impacts can
482 be investigated by comparing the historical and future wave power in the selected locations. Fig.
483 6 illustrates the spatial distribution of mean wave energy over the study area.

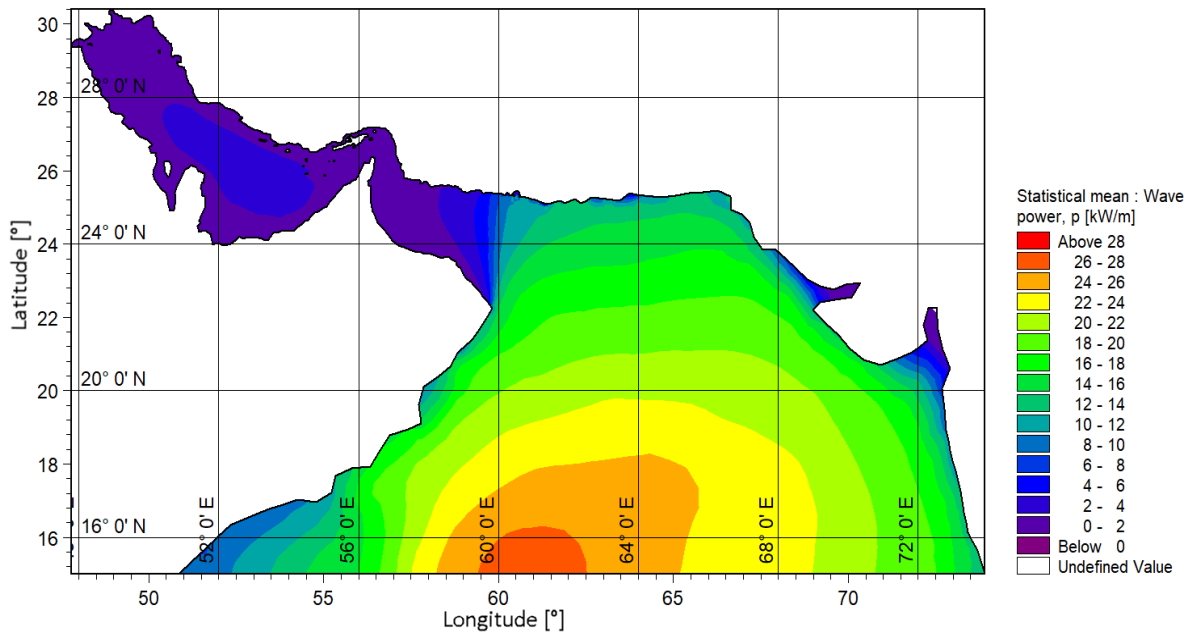
484

485

486



487 a)



488 b)

489 Fig. 6. Spatial distribution of mean wave energy for the a) historical (1981 to 2000) and, b)
490 **future periods (2081-2100)**

491 As observed from Fig. 6, the wave energy over the study area shows a great degree of variability
492 where the offshore and eastern parts of the gulf are prone to more energetic waves. Regarding the
493 future scenario, an increasing trend in wave energy can be observed over the study area. Four

494 locations along the northern part of the gulf have been selected to provide more in-depth analysis
 495 of energy variability. They were considered according to the spatial distribution of wave energy,
 496 distance from the coasts, vicinity to the industrial or residential regions, etc. Hence, the
 497 computations were carried out for the total and exploitable storages of wave power as presented in
 498 Table 6 for both historical and future periods.

499

500 Table 6. Historical and future computations of total and exploitable wave power

Climate	Parameter	Jask	Chabahar	Beris	Pasabandar
Hist.	P_{ave} (kW/m)	1.39	11.02	9.01	10.87
	E_t (kWh/m)	12186	96547	78934	95192
	t_e (hr)	396.2	1000.1	971.8	1011.3
	E_e (kWh/m)	551	11022	8757	10989
	<i>Exploitability</i> %	5%	11%	11%	12%
SSP5-8.5	P_{ave} (kW/m)	2.01	13.58	10.89	13.43
	E_t (kWh/m)	17633	118946	95390	117651
	t_e (hr)	448.4	1017.3	1008.2	1045.4
	E_e (kWh/m)	903	13813	10978	14040
	<i>Exploitability</i> %	5%	12%	12%	12%

501

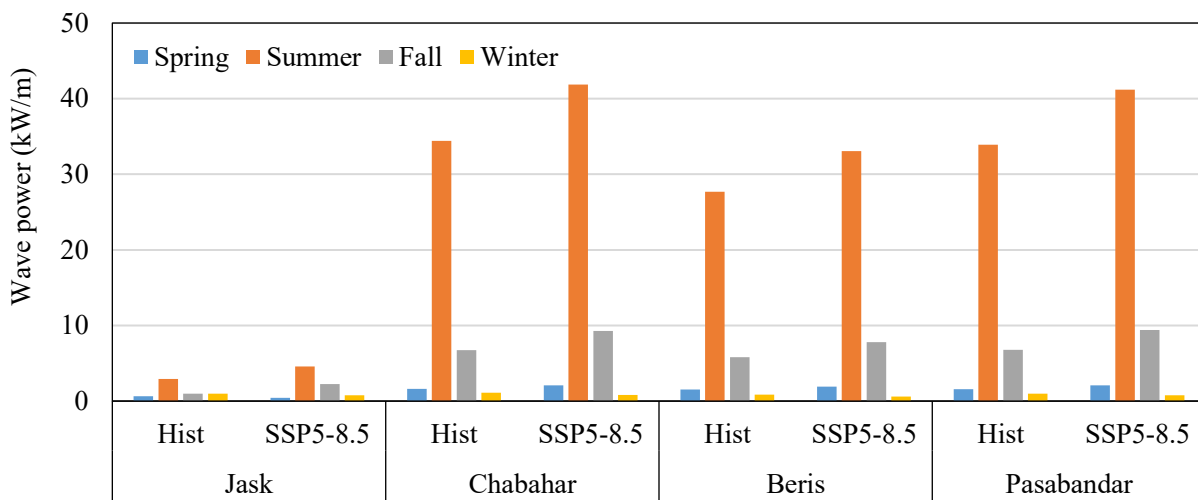
502 The results presented in Table 6 show an increasing trend in total and exploitable wave power
 503 under SSP5-8.5. An average increase of about 26% in exploitable energy for the three eastern sites
 504 is obtained under the future SSP. In addition, a remarkable difference can be detected in total and
 505 exploitable wave energy, where the exploitable energy can be about 12% of the total energy in
 506 some locations. In the other locations, it can be found as 5%. Considering mean annual values of
 507 wave power, it is expected that the wave power values are to be increased according to SSP5-8.5.
 508 Regarding the wave power variation for the different locations, it can be inferred that the eastern
 509 regions have higher potential than the western areas. For example, the wave power in Pasabandar
 510 has a magnitude of about 6.7 and 7.8 times the corresponding values in Jask for the historical and

511 the future periods, respectively. The seasonal distribution of wave energy can provide additional
 512 details to the wave energy assessment considering the variability of energy demand and supply.

513

514 3.4. Seasonal and monthly variability of historical and future wave power

515 Temporal variation of wave energy is required to investigate the stability of resources in the
 516 short term. Fig. 7 depicts the seasonal variability of wave power for the different locations in the
 517 historical and future periods.



518

519 Fig 7. Seasonal variability of wave power for the historical (1981-2000) and future (2081-2100)
 520 periods

521

522 According to Fig. 7, considerable seasonal variability in wave power can be seen. Generally, in all
 523 the stations, the highest energy exists in the summer, while fall ranks the second season in terms
 524 of the highest wave power. The pattern of public circulation in this region is significantly
 525 influenced by summer and winter monsoons driven by the latent temperature difference between
 526 land and sea. The Winter monsoon occurs from November to April and is accompanied by
 527 northeast winds with an average speed of below 5 meters per second. Summer monsoon (July-
 528 September) causes energetic and continuous winds in the south and southwest with an average

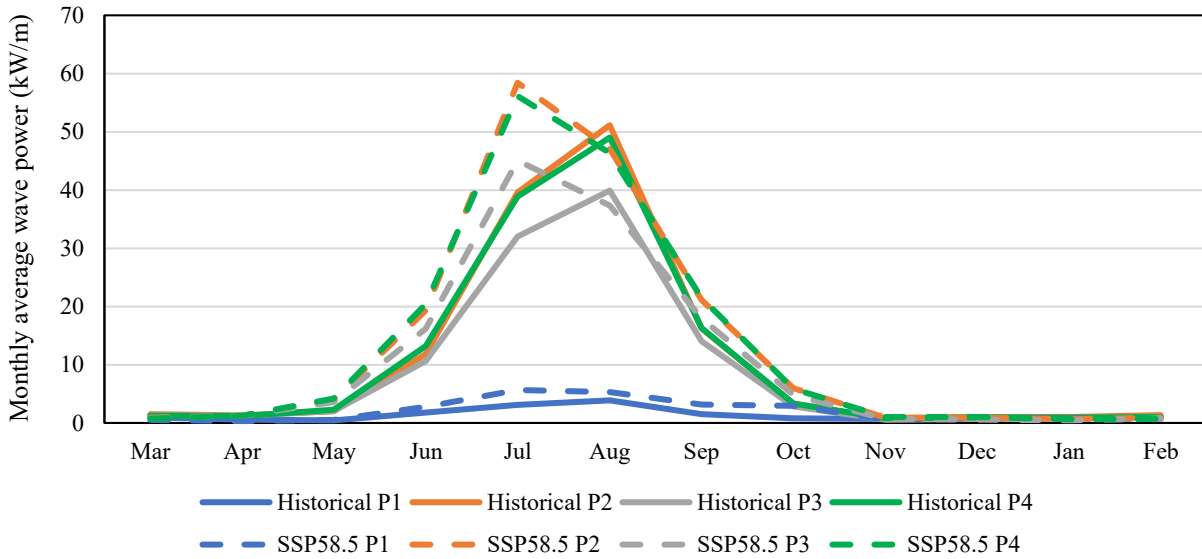
529 speed of 15 meters per second. During the summer, this constant wind regime significantly
530 contributes to ocean circulation and biochemical processes in the northern Arabian Sea and the
531 Gulf of Oman. The remaining months, either after the monsoon (October) or before the monsoon
532 (May and June), are considered with a lesser transition phase and stable wind pattern. Analysis of
533 annual wind data from the Gulf of Oman and the northern Arabian Sea shows that the predominant
534 wind direction in the Gulf of Oman is from west to northwest, while in the Arabian Sea, the
535 southwest direction is dominant [51].

536 As it can be seen in Fig. 7, the wave energy increases in this scenario. In winter, on the other hand,
537 the wave energy decreases in reverse. This somehow confirms that the Monsoon cycle can be
538 intensified. According to seasonal variability assessment, wave power increases in SPP5-8.5
539 scenario in spring, summer, and fall by about 13%, 30%, and 59%, respectively. However, it
540 decreases about 24 percent in winter on average in four stations.

541 During the summertime in the northern hemisphere, summer monsoons blow southwest and
542 generate strong, consistent waves [52]. During the summer monsoons, the western coasts of India
543 experience wave heights up to 6 *m*, whereas in other seasons, the significant wave height is about
544 1.5 *m* [53]. Similar trends have been demonstrated in the previous studies outlining wind and wave
545 climate in the Gulf of Oman [51,54]. On the contrary, the wave power reaches the lowest values
546 during winter and spring. This finding is promising to have higher wave energy in summer when
547 the demand for the energy supply increases due to hot weather conditions in summers. In this
548 regard, it is worth mentioning that the energy consumption for indoor space cooling (air
549 conditioning) has a large share of energy consumption in the northern land of the study area.
550 Therefore, an excess demand for energy in summer is possible due to the higher temperatures.
551 Hence, wave energy as a supplementary energy supply can be considered.

552 The future wave power under SSP5-8.5 is higher than the historical values indicating an increase
553 of 21-45% in the future. *P4* and *P2* have relatively higher wave energy than the other stations.
554 Since the seasonal variability is considerable in the study area, assessing intra-annual variation on
555 a monthly scale seems necessary. Hence, the monthly variation of the wave energy is illustrated in
556 Fig. 8.

557



558

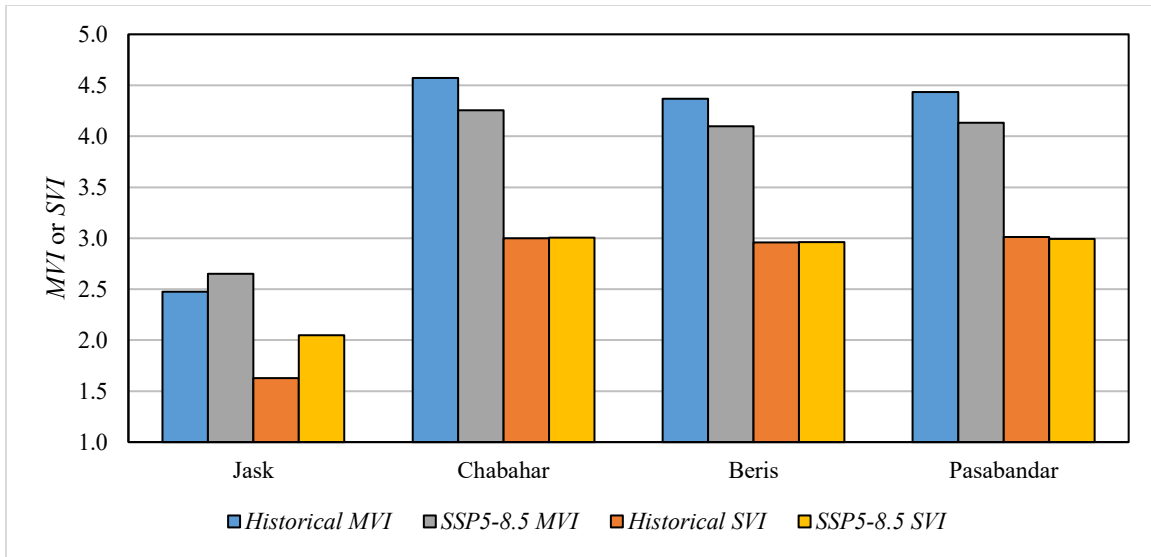
559 Fig 8. Monthly variation of wave power in the selected stations for the historical (1981-2000)
 560 and future (2081-2100) periods

561

562 The monthly variability of wave power (Fig. 8) reveals that July and August are among the highest
 563 energetic months, while the months in winter and fall have the lowest energy potential. In general,
 564 a high variation of wave power in different months can be detected, which is not desirable in terms
 565 of the sustainability of energy on the intra-annual scale. Comparison of the results of the historical
 566 and future periods shows that the future projections have higher values than historical ones. In
 567 order to quantitatively assess the intra-annual variability, historical and future *MVI* and *SVI* values
 568 are shown in Fig. 9.

569

570



571

572 Fig. 9. *MVI* and *SVI* values for the historical (1981-2000) and future (2081-2100) periods

573

574 Regarding Fig. 9, Jask has the most stable conditions in terms of monthly and seasonal variability,
 575 with the lowest *MVI* and *SVI* values for both historical and future periods. Other stations have
 576 higher *MVI* and *SVI*, which is not in favor of wave power stability. However, in Pasabandar, which
 577 wave power is remarkable, the variability indices are desirable to guarantee sustainability and
 578 availability of the power during different seasons. Since the high values for these indices are
 579 initiated by extremely high power in summer, it can be promising due to the increasing demand
 580 for electricity in summer when hot weather and more activities are expected. Therefore, the wave
 581 power may be considered an opportunity to compensate for the summer energy shortage for the
 582 study area. Finally, accessibility and availability of the wave power under the future high emission
 583 scenario are needed to be considered for operational purposes. More details have been discussed
 584 in the following sub-section.

585

586 3.4. Accessibility and availability of wave power under SSP5-8.5

587 Part of the assessment of wave energy is defined by the accessibility for operation and maintenance
 588 of the installations (e.g., Wave Energy Convertors (WECs)). Along with accessibility, the
 589 availability of the extractable wave energy is another sea state specification that is required to be

590 taken into account. The criterion specifying the sea state conditions is considered in terms of
 591 significant wave height. Moreover, several extreme events may yield high computational wave
 592 power, but not in favor of the operation of WECs. This is investigated as the percentage of the
 593 hours lying in cut-in and cut-off wave heights to the total hours. In this study, cut-in and cut-off
 594 values for significant wave height are 0.5 m and 4 m, respectively [50]. Table 7 represents the
 595 accessibility and availability values in different locations and future periods (SSP5-8.5).

596

597 Table 7. Annual and seasonal accessibility and availability of the wave characteristics in each
 598 station

Station	Season	Accessibility %							Availability %
		$H_s < 1.0$	$H_s < 1.5$	$H_s < 2.0$	$H_s < 2.5$	$H_s < 3.0$	$H_s < 3.5$	$H_s < 4.0$	$0.5 < H_s < 4.0$
Jask	Total	79.8	96.0	99.1	99.7	99.9	100.0	100.0	62.8
	Spring	91.1	97.9	99.5	99.8	99.9	100.0	100.0	51.2
	Summer	49.0	92.1	99.3	100.0	100.0	100.0	100.0	99.1
	Fall	91.6	99.4	100.0	100.0	100.0	100.0	100.0	52.0
	Winter	88.0	94.8	97.8	99.1	99.7	100.0	100.0	48.5
Chabahar	Total	51.7	67.5	76.9	84.9	93.2	99.1	100.0	81.0
	Spring	62.3	88.1	98.1	99.4	100.0	100.0	100.0	84.3
	Summer	0.2	5.5	16.2	42.2	73.3	96.5	99.9	99.9
	Fall	57.5	81.9	96.1	99.3	100.0	100.0	100.0	83.5
	Winter	87.7	95.3	97.9	99.2	99.8	100.0	100.0	55.8
Beris	Total	55.1	71.1	79.8	87.8	96.3	99.9	100.0	76.3
	Spring	66.2	93.2	99.2	99.9	100.0	100.0	100.0	84.7
	Summer	0.5	6.7	24.0	52.3	85.4	99.6	100.0	100.0
	Fall	62.8	88.1	97.3	99.7	100.0	100.0	100.0	79.3
	Winter	91.6	97.2	99.1	99.8	100.0	100.0	100.0	40.3
Pasabandar	Total	52.0	67.2	76.4	84.3	92.1	98.7	100.0	81.2
	Spring	61.8	88.3	99.4	100.0	100.0	100.0	100.0	85.0
	Summer	0.1	3.9	12.5	38.9	68.8	94.9	99.9	99.9
	Fall	56.6	80.1	95.4	99.3	100.0	100.0	100.0	85.1
	Winter	90.2	97.3	98.9	99.5	99.8	100.0	100.0	54.2

599

600 According to Table 7, it can be found that for Jask, accessibility is higher than in other locations.
 601 Generally, all the locations have accessibility higher than 50%, which is suitable to meet the
 602 operation and maintenance of the equipment. These values for significant wave heights of less than

603 1 m reach 80% or higher during winter, while they decrease under 50% in summer. However, for
604 larger waves but lower than 3.5 or 4 m, the accessibility for all locations and seasons converges to
605 100%. In terms of availability, Pasabandar and Chabahar have the highest availability, indicating
606 a better condition for sustainable wave energy extraction. For Jask and Beris, the availability factor
607 is still higher than 60%. The highest and the lowest availability of energy resources are related to
608 the summer and winter seasons, respectively. Seasonal and total computations of accessibility and
609 availability factors for different locations demonstrate the efficiency of the wave energy extraction
610 for the region.

611 Fig. 9 shows that with the increase in the availability factor, accessibility usually decreases. The
612 highest accessibility and availability values in different stations belong to Jask in the west and
613 Pasabandar in the east, respectively. Therefore, the point in the western part of the region is more
614 suitable in terms of accessibility, while the eastern parts are more suitable for availability. Hence,
615 it can be found that moving from the western part of the region to the eastern parts, availability
616 increases slightly while accessibility decreases. Considering both factors, the energy potential in
617 all stations is considered satisfactory where accessibility and availability are higher than 50%.

618 4. Discussion

619

620 5. Conclusions

621 As a source of renewable energy, wave power plays an important role in supplying the increasing
622 electricity demand. Different factors are investigated to define the suitability of a station for wave
623 energy exploitation, including the amount of wave power, exploitable storage of wave energy,
624 availability, accessibility, and the intra-annual fluctuation of the mentioned parameters. Moreover,
625 climate change may remarkably influence the wave energy potential since it is influenced by wind
626 patterns and, consequently, the wave climate. This study employed near-surface wind field
627 simulations from a CMIP6 model, i.e., CNRM, for the historical period and a high emission future
628 projection (SSP5-8.5) to drive the wave model and generate the wave characteristics for two 20-
629 yearly periods (1981- 2000 and 2018-2100). Subsequently, wave power for different locations in
630 the northern strip of the Gulf of Oman was calculated. Before performing the numerical wave

631 modeling, a new statistical bias-correction technique based on the Weibull distribution was
 632 proposed to modify CNRM wind field compared with ERA-5 dataset. The main findings of this
 633 study can be summarized as follow.

- 634 • The proposed bias-correction technique partially improved the performance of the existing
 635 bias-correction method in additive form. It was found that using the multiplicative form to
 636 modify Weibull parameters can be considered as an alternative to the additive form.
- 637 • Wave power is drastically higher in summer than in other seasons, indicating a significant
 638 temporal and seasonal variability in wave power distribution.
- 639 • Under a high emission scenario, an increase in a range of 21-45% in the annual wave power
 640 is estimated in the study area. The relative intensification of wave power is greater in the
 641 western part (Jask station).
- 642 • Under a high emission scenario, wave energy increases in spring, summer, and fall by 13%,
 643 30%, and 59 %, respectively. Conversely, wave power decreases about 24% in winter when
 644 averaged over the four stations.
- 645 • All the stations have availability and accessibility factors higher than 50%, demonstrating
 646 the suitability of the stations for wave energy extraction and wave energy converter
 647 installation. According to the results, accessibility increases 4 % in SSP5-8.5.
- 648 • Exploitable wave energy increases as total wave energy is between 15 and 23 %, depending
 649 on the station.

650 This study demonstrated a high spatio-temporal variability with an increasing trend in wave power
 651 under a future warmer climate. However, summer is the season with the highest energy, which is
 652 promising due to the increasing energy demand. Moreover, it was revealed that wave power will
 653 increase slightly in the future. The exploitable storage of wave energy exceeds 14MWh in the
 654 study area. Results of this study indicating the suitable potential of wave energy and its
 655 sustainability can provide useful information for wave energy developers to consider this type of
 656 renewable energy for rural and urban regions nearby the gulf. The growing energy demand may
 657 lead to joint wind-wave energy (eventually also solar) resource exploitation, increasing the
 658 competitiveness of ocean energy.

659

660 Acknowledgment

661 The authors are thankful to the DHI Group for providing us with the student license of MIKE 21
662 SW. B.K. has been supported by the Hakubi Center for Advanced Research at Kyoto University,
663 and JSPS Grants-in-Aid for Scientific Research (KAKENHI), grant No. 20K04705, supported by
664 the Ministry of Education, Culture, Sports, Science, and Technology of Japan (MEXT).

665

666 References

- 667 [1] M. Şan, A. Akpınar, B. Bingölbalı, M. Kankal, Geo-spatial multi-criteria evaluation of
668 wave energy exploitation in a semi-enclosed sea, *Energy*. (2021).
669 <https://doi.org/10.1016/j.energy.2020.118997>.
- 670 [2] J. Portilla, J. Sosa, L. Cavaleri, Wave energy resources: Wave climate and exploitation,
671 *Renew. Energy*. (2013). <https://doi.org/10.1016/j.renene.2013.02.032>.
- 672 [3] A.G. Majidi, B. Bingölbalı, A. Akpınar, E. Rusu, Wave power performance of wave
673 energy converters at high-energy areas of a semi-enclosed sea, *Energy*. 220 (2021)
674 119705. <https://doi.org/https://doi.org/10.1016/j.energy.2020.119705>.
- 675 [4] A. Ulazia, G. Esnaola, P. Serras, M. Penalba, On the impact of long-term wave trends on
676 the geometry optimisation of oscillating water column wave energy converters, *Energy*.
677 (2020). <https://doi.org/10.1016/j.energy.2020.118146>.
- 678 [5] M.J. Alizadeh, T. Alinejad-Tabrizi, M.R. Kavianpour, S. Shamshirband, Projection of
679 spatiotemporal variability of wave power in the Persian Gulf by the end of 21st century:
680 GCM and CORDEX ensemble, *J. Clean. Prod.* (2020).
681 <https://doi.org/10.1016/j.jclepro.2020.120400>.
- 682 [6] J.P. Sierra, M. Casas-Prat, E. Campins, Impact of climate change on wave energy
683 resource: The case of Menorca (Spain), *Renew. Energy*. (2017).
684 <https://doi.org/10.1016/j.renene.2016.08.060>.
- 685 [7] L. Rusu, Evaluation of the near future wave energy resources in the Black Sea under two

- 686 climate scenarios, *Renew. Energy*. (2019). <https://doi.org/10.1016/j.renene.2019.04.092>.
- 687 [8] M.V.W. Cuttler, J.E. Hansen, R.J. Lowe, Seasonal and interannual variability of the wave
688 climate at a wave energy hotspot off the southwestern coast of Australia, *Renew. Energy*.
689 (2020). <https://doi.org/10.1016/j.renene.2019.08.058>.
- 690 [9] D. Iribarren, M. Martín-Gamboa, Z. Navas-Anguita, D. García-Gusano, J. Dufour,
691 Influence of climate change externalities on the sustainability-oriented prioritisation of
692 prospective energy scenarios, *Energy*. (2020).
693 <https://doi.org/10.1016/j.energy.2020.117179>.
- 694 [10] B.G. Reguero, I.J. Losada, F.J. Méndez, A recent increase in global wave power as a
695 consequence of oceanic warming, *Nat. Commun.* (2019). [https://doi.org/10.1038/s41467-](https://doi.org/10.1038/s41467-018-08066-0)
696 [018-08066-0](https://doi.org/10.1038/s41467-018-08066-0).
- 697 [11] A. Ulazia, M. Penalba, G. Ibarra-Berastegui, J. Ringwood, J. Saénz, Wave energy trends
698 over the Bay of Biscay and the consequences for wave energy converters, *Energy*. (2017).
699 <https://doi.org/10.1016/j.energy.2017.09.099>.
- 700 [12] A. Martinez, G. Iglesias, Wind resource evolution in Europe under different scenarios of
701 climate change characterised by the novel Shared Socioeconomic Pathways, *Energy*
702 *Convers. Manag.* 234 (2021) 113961.
703 <https://doi.org/https://doi.org/10.1016/j.enconman.2021.113961>.
- 704 [13] A. Martinez, G. Iglesias, Climate change impacts on wind energy resources in North
705 America based on the CMIP6 projections, *Sci. Total Environ.* 806 (2022) 150580.
706 <https://doi.org/https://doi.org/10.1016/j.scitotenv.2021.150580>.
- 707 [14] R. Le Roux, M. Katurji, P. Zawar-Reza, H. Quénol, A. Sturman, Comparison of statistical
708 and dynamical downscaling results from the WRF model, *Environ. Model. Softw.* (2018).
709 <https://doi.org/10.1016/j.envsoft.2017.11.002>.
- 710 [15] M.J. Alizadeh, M.R. Kavianpour, B. Kamranzad, A. Etemad-Shahidi, A Weibull
711 Distribution Based Technique for Downscaling of Climatic Wind Field, *Asia-Pacific J.*
712 *Atmos. Sci.* (2019). <https://doi.org/10.1007/s13143-019-00106-z>.

- 713 [16] B. Liu, K.B. Costa, L. Xie, F.H.M. Semazzi, Dynamical downscaling of climate change
 714 impacts on wind energy resources in the contiguous United States by using a limited-area
 715 model with scale-selective data assimilation, *Adv. Meteorol.* (2014).
 716 <https://doi.org/10.1155/2014/897246>.
- 717 [17] J. Ringwood, G. Brandle, A new world map for wave power with a focus on variability,
 718 in: *Proc. 11th Eur. Wave Tidal Energy Conf., European Wave and Tidal Energy*
 719 *Conference 2015*, 2015.
- 720 [18] R.G. Coe, S. Ahn, V.S. Neary, P.H. Kobos, G. Bacelli, Maybe less is more: Considering
 721 capacity factor, saturation, variability, and filtering effects of wave energy devices, *Appl.*
 722 *Energy.* (2021). <https://doi.org/10.1016/j.apenergy.2021.116763>.
- 723 [19] D. Khojasteh, D. Khojasteh, R. Kamali, A. Beyene, G. Iglesias, Assessment of renewable
 724 energy resources in Iran; with a focus on wave and tidal energy, *Renew. Sustain. Energy*
 725 *Rev.* 81 (2018) 2992–3005. <https://doi.org/https://doi.org/10.1016/j.rser.2017.06.110>.
- 726 [20] A. Saket, A. Etemad-Shahidi, Wave energy potential along the northern coasts of the Gulf
 727 of Oman, Iran, *Renew. Energy.* (2012). <https://doi.org/10.1016/j.renene.2011.09.024>.
- 728 [21] B. Kamranzad, V. Chegini, A. Etemad-Shahidi, Temporal-spatial variation of wave
 729 energy and nearshore hotspots in the Gulf of Oman based on locally generated wind
 730 waves, *Renew. Energy.* (2016). <https://doi.org/10.1016/j.renene.2016.03.084>.
- 731 [22] B. Kamranzad, S. Hadadpour, A multi-criteria approach for selection of wave energy
 732 converter/location, *Energy.* 204 (2020) 117924.
 733 <https://doi.org/https://doi.org/10.1016/j.energy.2020.117924>.
- 734 [23] DHI, MIKE 21 SW - Spectral Wave Model: Scientific Documentation, DHI Water
 735 Environ. (2014).
- 736 [24] A.S. Ribeiro, M. deCastro, X. Costoya, L. Rusu, J.M. Dias, M. Gomez-Gesteira, A Delphi
 737 method to classify wave energy resource for the 21st century: Application to the NW
 738 Iberian Peninsula, *Energy.* (2021). <https://doi.org/10.1016/j.energy.2021.121396>.
- 739 [25] H. Hersbach, B. Bell, P. Berrisford, S. Hirahara, A. Horányi, J. Muñoz-Sabater, J. Nicolas,

- 740 C. Peubey, R. Radu, D. Schepers, A. Simmons, C. Soci, S. Abdalla, X. Abellan, G.
 741 Balsamo, P. Bechtold, G. Biavati, J. Bidlot, M. Bonavita, G. De Chiara, P. Dahlgren, D.
 742 Dee, M. Diamantakis, R. Dragani, J. Flemming, R. Forbes, M. Fuentes, A. Geer, L.
 743 Haimberger, S. Healy, R.J. Hogan, E. Hólm, M. Janisková, S. Keeley, P. Laloyaux, P.
 744 Lopez, C. Lupu, G. Radnoti, P. de Rosnay, I. Rozum, F. Vamborg, S. Villaume, J.N.
 745 Thépaut, The ERA5 global reanalysis, Q. J. R. Meteorol. Soc. (2020).
 746 <https://doi.org/10.1002/qj.3803>.
- 747 [26] R. Seferian, CNRM-CERFACS CNRM-ESM2-1 model output prepared for CMIP6
 748 CMIP, (2018). <https://doi.org/10.22033/ESGF/CMIP6.1391>.
- 749 [27] K. Mahmoodi, H. Ghassemi, A. Razminia, Temporal and spatial characteristics of wave
 750 energy in the Persian Gulf based on the ERA5 reanalysis dataset, Energy. (2019).
 751 <https://doi.org/10.1016/j.energy.2019.115991>.
- 752 [28] M. Beyramzadeh, S.M. Siadatmousavi, M.H. Derkani, Calibration and skill assessment of
 753 two input and dissipation parameterizations in WAVEWATCH-III model forced with
 754 ERA5 winds with application to Persian Gulf and Gulf of Oman, Ocean Eng. (2021).
 755 <https://doi.org/10.1016/j.oceaneng.2020.108445>.
- 756 [29] M.J. Alizadeh, M.R. Kavianpour, B. Kamranzad, A. Etemad-Shahidi, A distributed wind
 757 downscaling technique for wave climate modeling under future scenarios, Ocean Model.
 758 145 (2020) 101513. <https://doi.org/10.1016/j.ocemod.2019.101513>.
- 759 [30] A. Patra, P.K. Bhaskaran, Temporal variability in wind–wave climate and its validation
 760 with ESSO-NIOT wave atlas for the head Bay of Bengal, Clim. Dyn. (2017).
 761 <https://doi.org/10.1007/s00382-016-3385-z>.
- 762 [31] M. Wandres, C. Pattiaratchi, Y. Hetzel, E.M.S. Wijeratne, The response of the southwest
 763 Western Australian wave climate to Indian Ocean climate variability, Clim. Dyn. (2018).
 764 <https://doi.org/10.1007/s00382-017-3704-z>.
- 765 [32] B.C. O’Neill, C. Tebaldi, D.P. Van Vuuren, V. Eyring, P. Friedlingstein, G. Hurtt, R.
 766 Knutti, E. Kriegler, J.F. Lamarque, J. Lowe, G.A. Meehl, R. Moss, K. Riahi, B.M.
 767 Sanderson, The Scenario Model Intercomparison Project (ScenarioMIP) for CMIP6,

- 768 Geosci. Model Dev. (2016). <https://doi.org/10.5194/gmd-9-3461-2016>.
- 769 [33] A. Lira-Loarca, F. Ferrari, A. Mazzino, G. Besio, Future wind and wave energy resources
770 and exploitability in the Mediterranean Sea by 2100, *Appl. Energy*. (2021).
771 <https://doi.org/10.1016/j.apenergy.2021.117492>.
- 772 [34] A. Ahmadalipour, A. Rana, H. Moradkhani, A. Sharma, Multi-criteria evaluation of
773 CMIP5 GCMs for climate change impact analysis, *Theor. Appl. Climatol.* (2017).
774 <https://doi.org/10.1007/s00704-015-1695-4>.
- 775 [35] A. Voldoire, D. Saint-Martin, S. Sénési, B. Decharme, A. Alias, M. Chevallier, J. Colin, J.
776 -F. Guérémy, M. Michou, M. -P. Moine, P. Nabat, R. Roebrig, D. y Mélia, R. Sférian, S.
777 Valcke, I. Beau, S. Belamari, S. Berthet, C. Cassou, J. Cattiaux, J. Deshayes, H. Douville,
778 C. Ethé, L. Franchistéguy, O. Geoffroy, C. Lévy, G. Madec, Y. Meurdesoif, R. Msadek,
779 A. Ribes, E. Sanchez-Gomez, L. Terray, R. Waldman, Evaluation of {CMIP6} {DECK}
780 {Experiments} {With} {CNRM}-{CM6}-1, *J. Adv. Model. Earth Syst.* 11 (2019) 2177–
781 2213. <https://doi.org/10.1029/2019MS001683>.
- 782 [36] A. Ouammi, R. Sacile, D. Zejli, A. Mimet, R. Benchrif, Sustainability of a wind power
783 plant: Application to different Moroccan sites, *Energy*. (2010).
784 <https://doi.org/10.1016/j.energy.2010.07.010>.
- 785 [37] I.A. Anton, L. Rusu, C. Anton, Nearshore wave dynamics at Mangalia beach simulated by
786 spectral models, *J. Mar. Sci. Eng.* (2019). <https://doi.org/10.3390/jmse7070206>.
- 787 [38] M.A. Hoque, W. Perrie, S.M. Solomon, Evaluation of two spectral wave models for wave
788 hindcasting in the Mackenzie Delta, *Appl. Ocean Res.* (2017).
789 <https://doi.org/10.1016/j.apor.2016.11.009>.
- 790 [39] M.H. Moeini, A. Etemad-Shahidi, Application of two numerical models for wave
791 hindcasting in Lake Erie, *Appl. Ocean Res.* (2007).
792 <https://doi.org/10.1016/j.apor.2007.10.001>.
- 793 [40] P.G. Remya, R. Kumar, S. Basu, A. Sarkar, Wave hindcast experiments in the Indian
794 Ocean using MIKE 21 SW model, *J. Earth Syst. Sci.* (2012).

- 795 <https://doi.org/10.1007/s12040-012-0169-7>.
- 796 [41] R.C. Ris, L.H. Holthuijsen, N. Booij, A third-generation wave model for coastal regions 2.
797 Verification, *J. Geophys. Res. Ocean.* (1999). <https://doi.org/10.1029/1998jc900123>.
- 798 [42] G.J. Komen, L. Cavaleri, M. Donelan, K. Hasselmann, S. Hasselmann, P.A.E.M. Janssen,
799 Dynamics and Modelling of Ocean Waves, 1994.
800 <https://doi.org/10.1017/cbo9780511628955>.
- 801 [43] M.J. Tucker, E.G. Pitt, Waves in ocean engineering, (2001).
802 <http://books.google.com/books?id=0v5RAAAAMAAJ>.
- 803 [44] R.E. and A.U.P. (U.S.), Technology white paper on wave energy potential on the U.S.
804 Outer Continental Shelf., (2006). <https://purl.fdlp.gov/GPO/LPS94821>.
- 805 [45] B. Kamranzad, S. Hadadpour, A multi-criteria approach for selection of wave energy
806 converter/location, *Energy*. 204 (2020) 117924.
807 <https://doi.org/https://doi.org/10.1016/j.energy.2020.117924>.
- 808 [46] L. Margheritini, D. Vicinanza, P. Frigaard, SSG wave energy converter: Design, reliability
809 and hydraulic performance of an innovative overtopping device, *Renew. Energy*. 34
810 (2009) 1371–1380.
- 811 [47] C.W. Zheng, C.Y. Li, X. Chen, J. Pan, Numerical Forecasting Experiment of the Wave
812 Energy Resource in the China Sea, *Adv. Meteorol.* (2016).
813 <https://doi.org/10.1155/2016/5692431>.
- 814 [48] C.W. Zheng, J. Pan, J.X. Li, Assessing the China Sea wind energy and wave energy
815 resources from 1988 to 2009, *Ocean Eng.* (2013).
816 <https://doi.org/10.1016/j.oceaneng.2013.03.006>.
- 817 [49] M. Penalba, A. Ulazia, G. Ibarra-Berastegui, J. Ringwood, J. Sáenz, Wave energy
818 resource variation off the west coast of Ireland and its impact on realistic wave energy
819 converters' power absorption, *Appl. Energy*. (2018).
820 <https://doi.org/10.1016/j.apenergy.2018.04.121>.
- 821 [50] G. Lavidas, A. Agarwal, V. Venugopal, Availability and Accessibility for Offshore

- 822 Operations in the Mediterranean Sea, *J. Waterw. Port, Coastal, Ocean Eng.* (2018).
823 [https://doi.org/10.1061/\(asce\)ww.1943-5460.0000467](https://doi.org/10.1061/(asce)ww.1943-5460.0000467).
- 824 [51] N. Chaichitehrani, M.N. Allahdadi, Overview of Wind Climatology for the Gulf of Oman
825 and the Northern Arabian Sea, *Am. J. Fluid Dyn.* (2018).
- 826 [52] M. Seemanth, S.A. Bhowmick, R. Kumar, R. Sharma, Sensitivity analysis of dissipation
827 parameterizations in a third-generation spectral wave model, WAVEWATCH III for
828 Indian Ocean, *Ocean Eng.* (2016). <https://doi.org/10.1016/j.oceaneng.2016.07.023>.
- 829 [53] P. Vethamony, K. Sudheesh, S.P. Rupali, M.T. Babu, S. Jayakumar, A.K. Saran, S.K.
830 Basu, R. Kumar, A. Sarkar, Wave modelling for the north Indian Ocean using MSMR
831 analysed winds, *Int. J. Remote Sens.* 27 (2006) 3767–3780.
832 <https://doi.org/10.1080/01431160600675820>.
- 833 [54] V. Sanil Kumar, J. Singh, P. Pednekar, R. Gowthaman, Waves in the nearshore waters of
834 northern Arabian Sea during the summer monsoon, *Ocean Eng.* (2011).
835 <https://doi.org/10.1016/j.oceaneng.2010.11.009>.
- 836