

45 migration of CO₂ or brines from the CO₂ store may beneficially relieve reservoir fluid pressure (Cihan et al.,
46 2013). However unintended leakage of CO₂ or formation fluids would effect a number of stakeholders,
47 incurring economic and financial costs (Bielicki et al., 2014), environmental impact (Jones et al., 2015) and
48 also challenge the social and political acceptability of the technology (Ha-Duong and Loisel, 2009). As such
49 any incidence of leakage from engineered stores could have ramifications for the CCS industry on a global
50 scale, and so the viability of CCS depends on the reliable containment of injected CO₂ in the subsurface.

51 Legislation and guidelines developed for CCS set performance requirements that seek to minimize the risk of
52 leakage from the storage complex. The IPCC (2005) recommend that CO₂ stores should operate with less
53 than 1% CO₂ loss to the surface over 1,000 years. The US Department of Energy (US DOE) aims for 99%
54 containment of CO₂ injected for the purpose of geological storage (Bielicki et al., 2015), whereas the EU CCS
55 Directive (EU, 2009) requires CO₂ to remain 'permanently' in the storage formation. Any CO₂ that leaks from
56 the storage formation must therefore be quantified for reasons of performance assurance, as well as carbon
57 accounting (Dixon et al., 2015). Furthermore, legislation permitting subterranean CO₂ storage in the US, EU
58 and Japan or seabed in the North Atlantic require appropriate assessment of the risk of CO₂ leakage from
59 the intended storage reservoir, the potential impacts of CO₂ leakage on the environment, and means of
60 monitoring for leakage (Jones et al., 2015). In this context, 'environment' includes the near-subsurface (such
61 as underground sources of drinking water) or surface (terrestrial or marine) ecosystems, including human
62 health. Environmental impacts might source from the CO₂ itself (free phase or dissolved) and any co-injected
63 impurities, or brines displaced as a result of pressure perturbation from CO₂ injection and migration or
64 degraded by geochemical interaction of CO₂ and the surrounding rock, or mobilization of other fluids (e.g.
65 methane). In the event of leakage of CO₂ into the near surface environment, the leak must be mitigated and
66 the site eventually returned to baseline conditions.

67 To comply with site performance and monitoring requirements it is necessary to have a comprehensive
68 understanding of the possible spread and fate of CO₂ in the deep and shallow subsurface, and the potential
69 impacts of such leakage. Since approximately 40% of global storage capacity is located offshore (IEAGHG,
70 2008), it is important that leakage into marine and terrestrial environments is explored.

71 However, studies of sealing and non-sealing naturally occurring CO₂ reservoirs find that leakage to surface is
72 globally rare, even at sites that would not be considered suitably secure for CO₂ storage (Miocic et al., 2016,
73 Roberts et al., 2017b, Miocic et al., 2019). Should CO₂ migrate from the storage reservoir, multiple processes
74 will attenuate the CO₂, such that the likelihood of CO₂ reaching the surface is low (Alcalde et al., 2018, Roberts
75 et al., 2018). Improperly sealed wellbores present the most likely pathways of leakage (IPCC, 2005). There
76 has been no leakage to surface at pilot and commercial scale CO₂ storage operations to date. In the absence
77 of cases of CO₂ leakage, the CCS community have looked to natural analogues and field experiments to
78 further scientific understanding of CO₂ leakage and its impacts, and to develop monitoring approaches that
79 are capable of enabling any CO₂ leaks to be identified, attributed, and quantified (referred to as monitoring,
80 measurement and verification, MMV, techniques). Indeed, since 2006 a number of field-scale controlled
81 release experiments have been conducted around the world. The experiments release free phase or
82 dissolved CO₂ into the shallow subsurface to artificially simulate a CO₂ leak into the near-surface
83 environment. The experiments differ in regard to the geological and surface environments and experimental
84 set-up, including the injection rate and monitoring strategy. Since CO₂ release is controlled, and so the
85 injection rate and quantities are known, these experiments provide excellent opportunities to test methods
86 of measuring and quantifying CO₂ fate, and compare changes to environmental conditions and ecosystem
87 health. MMV methods for commercial-scale applications can be tested and calibrated while capability and
88 expertise is developed through learning-by-doing at the field site. Further, field experience of acquiring
89 baseline and post-release information is valuable for developing site monitoring protocols that fulfil the
90 decrees of current environmental legislation.

91 A number of recent reviews have excellently summarised the significant contribution that these experiments
92 have made to current scientific understanding of environmental impacts and state of the art monitoring
93 techniques (Feitz et al., 2014b, Jenkins et al., 2015, Jones et al., 2015, Lee et al., 2016). However to date there
94 has been no comprehensive examination of the collective learning at these sites with regards to the fate and
95 spread of the CO₂ and the surface manifestation of the leakage. Similarly, there have been no syntheses of
96 the injection rates and the leakage pathways that these sites mimic, nor a consolidation of lessons learned

97 for the design of an advanced experimental approach. To this end, we have collated a global dataset of field-
 98 scale controlled release experiments, detailing the experimental approach and findings about the surface
 99 and subsurface manifestation of the CO₂ release, the fate of the CO₂ and leakage quantification. The results
 100 were compared to observations from natural analogue and modelling studies and scrutinised to elucidate
 101 collective learnings and highlight where uncertainties remain. Future release experimental design and
 102 reporting will benefit from this work as the scientific community continue to seek methods to best
 103 characterise and monitor storage sites most effectively.

104 2. Method and approach: collating global CO₂ release experiments

105 We compiled a dataset of field-scale shallow controlled CO₂ release experiments for which the research
 106 results are publicly available (prior to May 2017). We focussed on field experiments that injected/released
 107 CO₂ into the subsurface with the aim that it would migrate and reach the surface or shallow subsurface rather
 108 than remain trapped in the injection formation.

109 Dataset variables, listed in Table 1, were populated through detailed review of published literature
 110 complemented by personal communication/interview with some of the key research scientists involved in
 111 projects or site custodians.

112 There is no standard unit for reporting CO₂ release rates into the subsurface and CO₂ fluxes from land surface
 113 or seabed to atmosphere or water column. As such, in the experiments we reviewed, rate of CO₂ leakage
 114 could be expressed in terms of mass (g, kg, tonnes) or volume (mL, L) or concentration (mol, mmol) per unit
 115 of time (which might be expressed as per second, per min, per hour, per day, per year). CO₂ flux, by definition,
 116 should be expressed as the rate of CO₂ leaked per unit area (usually m²). If no area unit was provided, the
 117 reported value is the CO₂ leakage rate (rate of CO₂ leaked), rather than flux specifically. To enable direct data
 118 comparison, where possible, we harmonised the CO₂ release rates and CO₂ fluxes so that dataset parameters
 119 were presented in standardised units (see Table 1). We elected to express CO₂ flux as g(CO₂)s⁻¹m⁻² and total
 120 rate of CO₂ leakage as g(CO₂)s⁻¹, but we also consider CO₂ leakage rate as tonnes per annum, t(CO₂)pa, since
 121 this is the standard unit for carbon accounting. If specific information or data were not reported or available,
 122 where possible, these values were inferred, calculated or estimated from the published information; for
 123 example, seep width could be inferred from the spatial distribution of CO₂ flux, or vertical leak velocity could
 124 be calculated from the information about CO₂ injection depth and surface arrival time. When converting, for
 125 example, from CO₂ volume to CO₂ mass, in the absence of specific temperature and pressure conditions at
 126 the site we assume CO₂ properties at STP.

127 We also noted the key monitoring tools that were used at each experiment - particularly for leakage
 128 quantification, CO₂ attribution and fate, including the presence of any added chemical tracers. The Table 1
 129 dataset is incomplete for many of the CO₂ release experiments studied; either those data were not collected
 130 or they are not yet publicly available.

131

132

Variable	Sub variable	Description	Units/note	SI
Basic descriptive information	Acronym	Acronym of the site name		Y
	Location	Longitude, latitude, country	degrees	Y
	Project aims	Principal research aims	broad aims	Y
		Release to surface intended or not		Y
	Funding body	Source of funding		N
		Funding total	€ Euros	Y
	Project partners			N
Key contact	Name and contact email of Principle Investigator		N	
Project status	Project completed / more CO ₂ releases intended		N	
Experiment set-up and site information	Well information	Depth below ground surface of CO ₂ injector	metres	Y
		Borehole type (deviated, vertical or inclined)		Y
		Properties of CO ₂ injector (single or multiple source)	Number of release points	Y
		Properties of the intended injection formation	Rock type; thickness	Y

	Geological information	Properties of the overlying formation	If different from injection formation	Y
		Soil properties	Type; thickness	Y
	Ecosystem type	Terrestrial/marine		Y
		The surface ecosystem at the site	Incl. vegetation type.	Y
Hydrological characteristics	Water table depth and flow direction	metres	Y	
CO₂ injection (for each experiment)	CO ₂ properties	CO ₂ source; gas composition	(% CO ₂)	Y
		δ ¹³ C composition	‰	Y
		Injected phase	Gas or water containing dissolved CO ₂	Y
	Injection rate	Steady, variable, or incremental		Y
		Maximum and minimum	g(CO ₂)/s ; gs ⁻¹	Y
	Injection periods	Date injection started & ceased	Day/month/year (time)	Y
		Total injection period	days	Y
	Quantity of CO ₂ injected	For each experiment	kg	Y
Overall	Number of experiments at the site		Y	
	Total quantity of CO ₂ injected	kg	Y	
CO₂ fate	Surface leakage	Surface leakage?	Y/N	
		Lag time (between start of injection and surface arrival)	hours	Y
		Flux rate*	g(CO ₂)/s/m ² ; gs ⁻¹ m ⁻²	Y
		Vertical leak velocity**	m/s	Y
		Proportion of injected CO ₂ released to surface	%	Y
		Distribution	Patchy, uniform	Y
		Hotspot radius***	metres	Y
		Leak location (with respect to the injector)	above injector / deviated	Y
	Temporal changes		Y	
	Subsurface spread (soil gas)	Detected in soil gas?	Y/N	
		Maximum soil gas concentration	%	Y
		Lag time (between start of injection and arrival in soil gas)	hours	Y
		Distribution		Y
	Subsurface spread (groundwater)	Detected in groundwater?	Y/N	
		Method / parameter (e.g. pH, conductivity, etc.)		
Lag time (between start of injection and detection in groundwater)		days		
Additional monitoring	Monitoring area	Area of surveillance	m ²	Y
	Baseline	Baseline monitoring period	days	Y
		What was monitored (e.g. CO ₂ flux, soil gas, plant/ecosystem diversity)		Y
	Recovery	Post-injection monitoring period	days	Y
Time to return to baseline conditions		Days	Y	

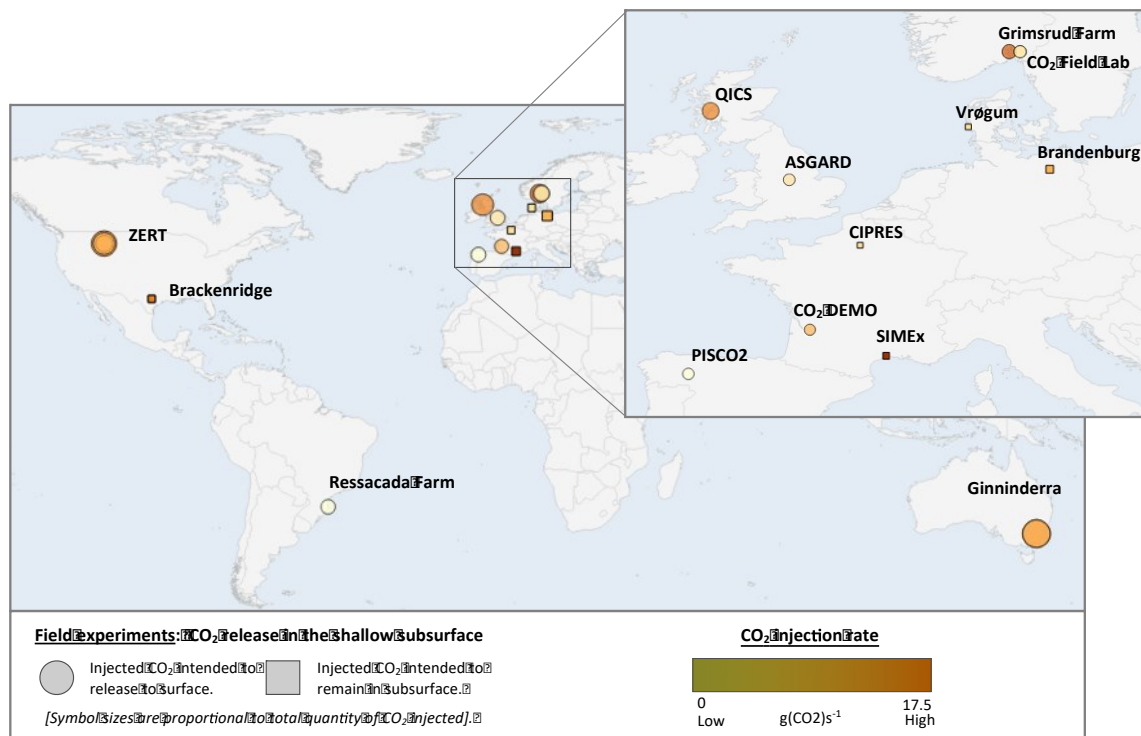
133 **Table 1.** Simplified list of variables collected for each CO₂ release experiment and the data included in the Supplementary
134 Information (SI). Data were collected from the published literature and from correspondence with project researchers.
135 Table 2 summarises key variables. *Where possible, if CO₂ flux (rate of CO₂ leakage per unit area) was not reported it
136 was calculated from information about the leakage area and the reported total leakage rate; **Vertical leak velocity
137 was calculated from the injection depth and the lag time between CO₂ injection and arrival at the surface (or near-
138 surface); ***Where information about the hotspot width was not reported, if possible, it was estimated from published
139 flux maps.

140

141 **3. Results and discussion**

142 The detailed dataset includes 14 different field experiment locations around the world. We refer to each by
 143 their project acronym or their location. These are shown in Figure 1, and a summary table of experimental
 144 parameters and results is provided in Table 2. Full data are available in the Supplementary Information.

145 We applied an injection depth cut off to the reviewed sites in order to focus on experiments which intended
 146 for the injected CO₂ to leak towards the surface. Only two experiments injected CO₂ deeper than 25 m; Plant
 147 Daniel (Trautz et al., 2012) and Cranfield (Yang et al., 2013) (which inject at 54 and 73 m, respectively) both
 148 of which were push-pull experiments investigating the effect of CO₂-bearing brines on groundwater quality.
 149 These were excluded from our dataset. The K-KOSEM EIT site in Korea is not included because at the time of
 150 writing there was little publicly available information about the experimental approach and results.



151
 152 *Fig. 1: Map of CO₂ release experiments around the world considered in our study. Symbols are coloured according to the*
 153 *CO₂ injection rates at the site, and sized proportional to the total amount of CO₂ injected over the lifetime of the sites*
 154 *(i.e. for some sites the total amount will be the sum of multiple injections).*

155
 156 *Table 2. Global compilation of controlled CO₂ release experiments at depths shallower than 25 m (with results published*
 157 *prior to May 2017). Half of the sites have conducted more than one release experiment ('No. of exp'), some of them*
 158 *preliminary tests, and the experiment phases often differed in the length ('inj. length (days)') and rate of CO₂ injection*
 159 *('Max inj. rate t(CO₂)pa'). The style of the injection also varies between sites, injecting CO₂ as a gas (g) or dissolved in*
 160 *water (diss), and via an inclined, vertical (v) or horizontal (h) well, and at a steady, incremental (incr.) or variable (var.)*
 161 *injection rate. For the CO₂ flux, modelled and measured are differentiated to make clear where values are informed by*
 162 *field measurements and where values are informed by modelling of the process based on other measured parameters.*

163 164 3.1 An overview of CO₂ release experiments

165 3.1.1 Global distribution

166 As can be seen in Figure 1, ten of the 14 field experiments included in this dataset are located in Europe
 167 (ASGARD, QICS, CO₂ Field Lab, Grimsrud Farm, Vrøgum, CO₂-Vadose/DEMO, CIPRES, SIMEx, Brandenburg,

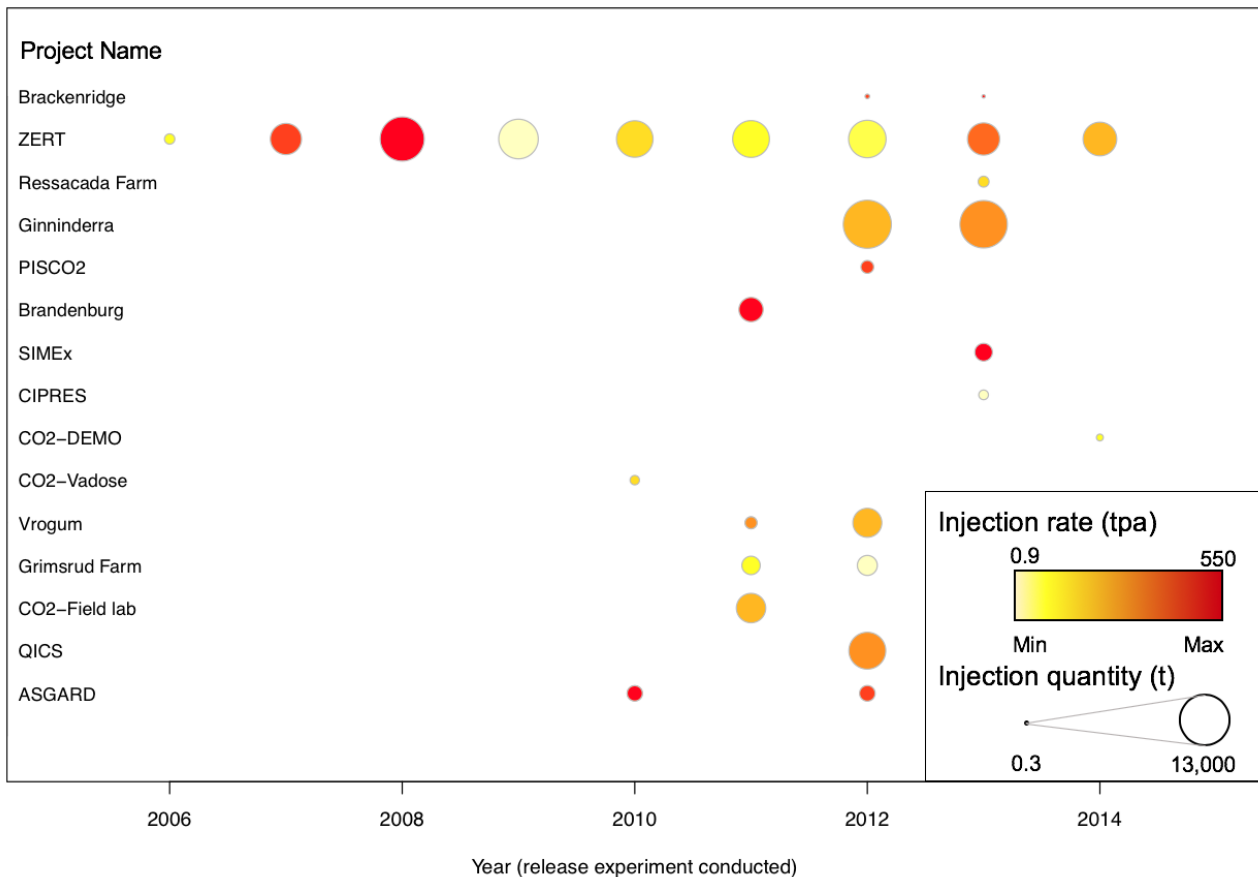
168 and PISCO2), with the remaining four located in Australia (Ginninderra), South America (Ressacada Farm)
 169 and the USA (ZERT, Brackenridge). Of all these field experiments to date, there has been only one subseabed
 170 CO₂ release; QICS. The CO₂-DEMO and CO₂-Vadose projects occur at the same location; the CO₂-Vadose
 171 project intended for CO₂ to migrate to surface, but was prevented from doing so by a thin clay layer (Cohen
 172 et al., 2013, Rillard et al., 2015), and so the experiment was modified to inject above this clay later as the
 173 CO₂-DEMO project.

174 3.1.2 Timeline of experiments

175 The experiments in our dataset were conducted between 2006-2014, and, as shown in Figure 2, most CO₂
 176 releases were conducted in the period 2011-13. Half of the projects performed more than one release
 177 experiment, and often the length and rate of CO₂ release differed between experiments (see Table 2). For
 178 example, release experiments have been conducted typically each summer at ZERT since 2007, where the
 179 injection rate has ranged from 0.62 to 3.47 gs⁻¹ (or 19 to 110 t(CO₂)pa) (Spangler et al., 2010).

180 In total, there have been 42 different CO₂ release experiments completed at the 14 sites that we reviewed,
 181 releasing a total of 82.8 t(CO₂) into the subsurface over 994 days (i.e. 2.7 years). This is not a complete
 182 dataset, since some preliminary experiments may not have been published, and some experiments may not
 183 yet be reported.

184 While a couple of projects are awaiting funding for additional experiments, at least half of the sites have
 185 completed their intended activities and thus the infrastructure has been dismantled.



186
 187 *Fig. 2: The quantity of CO₂ injected (circle size) and injection rate (colour) for each of the experiments conducted at the*
 188 *shallow CO₂ release projects investigated, and the year that the experiments were conducted. The ZERT site has*
 189 *conducted the most experiments, and also the most varied injection rate (deeper colours indicate higher injection rate).*

190

191 3.1.3 Objectives of CO₂ release experiments

192 The reviewed projects typically endeavoured to address one or more of the following principle aims:

- 193 1. Investigate **ecosystem** responses to the injected CO₂.
 194 2. Establish the **fluxes**, transformations and fate of CO₂ as it migrates from the injection point.
 195 3. Investigate **geochemical** interactions between CO₂ and groundwater.
 196 4. Test and calibrate **models** of CO₂ flow and fate.
 197 5. Test a broad or specific suite of monitoring, measurement and verification (**MMV**) techniques.

198 The primary aims for each project are shown in Table 3. Most of the projects focussed on developing MMV
 199 approaches for detecting or quantifying CO₂ leaks, and these approaches have enabled additional questions
 200 to be explored, primarily about the environmental impact of leaked CO₂ (on groundwater and ecosystems).
 201 As such, the projects are typically made up of interdisciplinary teams composed of multiple partners (often
 202 at least one publicly funded partner), and total budgets tend towards €1 M or greater.

203

204 *Table 3: Common primary objectives of CO₂ release experiments, categorised according to the list above (Section 3.1.3).*
 205 *This table is not comprehensive; secondary objectives might have been addressed by certain experiments at the site or*
 206 *certain team members. *Consecutive experiments at ZERT focussed on different objectives.*

n	Acronym	Ecosystem	Fluxes	Geochem	Models	MMV
1	ASGARD	<input checked="" type="checkbox"/>	<input type="checkbox"/>	<input type="checkbox"/>	<input type="checkbox"/>	<input checked="" type="checkbox"/>
2	QICS	<input checked="" type="checkbox"/>	<input checked="" type="checkbox"/>	<input checked="" type="checkbox"/>	<input checked="" type="checkbox"/>	<input checked="" type="checkbox"/>
3	CO ₂ Field Lab	<input type="checkbox"/>	<input type="checkbox"/>	<input type="checkbox"/>	<input checked="" type="checkbox"/>	<input checked="" type="checkbox"/>
4	Grimsrud Farm	<input checked="" type="checkbox"/>	<input type="checkbox"/>	<input type="checkbox"/>	<input type="checkbox"/>	<input type="checkbox"/>
5	Vrøgum	<input type="checkbox"/>	<input type="checkbox"/>	<input checked="" type="checkbox"/>	<input type="checkbox"/>	<input checked="" type="checkbox"/>
6	CO2DEMO	<input type="checkbox"/>	<input type="checkbox"/>	<input type="checkbox"/>	<input type="checkbox"/>	<input checked="" type="checkbox"/>
7	CIPRES	<input type="checkbox"/>	<input type="checkbox"/>	<input checked="" type="checkbox"/>	<input type="checkbox"/>	<input checked="" type="checkbox"/>
8	SIMEx	<input type="checkbox"/>	<input type="checkbox"/>	<input type="checkbox"/>	<input type="checkbox"/>	<input checked="" type="checkbox"/>
9	Brandenburg	<input type="checkbox"/>	<input type="checkbox"/>	<input checked="" type="checkbox"/>	<input type="checkbox"/>	<input checked="" type="checkbox"/>
10	PISCO2	<input checked="" type="checkbox"/>	<input type="checkbox"/>	<input type="checkbox"/>	<input checked="" type="checkbox"/>	<input checked="" type="checkbox"/>
11	Ginninderra	<input checked="" type="checkbox"/>	<input checked="" type="checkbox"/>	<input type="checkbox"/>	<input checked="" type="checkbox"/>	<input checked="" type="checkbox"/>
12	Ressacada Farm	<input type="checkbox"/>	<input type="checkbox"/>	<input type="checkbox"/>	<input type="checkbox"/>	<input checked="" type="checkbox"/>
13	ZERT*	<input checked="" type="checkbox"/>	<input checked="" type="checkbox"/>	<input checked="" type="checkbox"/>	<input type="checkbox"/>	<input checked="" type="checkbox"/>
14	Brackenridge	<input type="checkbox"/>	<input type="checkbox"/>	<input checked="" type="checkbox"/>	<input type="checkbox"/>	<input type="checkbox"/>
Total		6	3	6	3	12

207

208 The majority of the CO₂ release experiments sought to mimic the effects of leakage of CO₂ gas, either from a
 209 point source or vertical feature (such as from poorly sealed well casing), or a linear feature (such as from a
 210 fault). Two sites, CIPRES and Brackenridge, performed push-pull experiments using CO₂ dissolved in water to
 211 explore the effect of CO₂ on groundwater quality. Groundwater was pumped from the aquifer and saturated
 212 with CO₂ before being re-injected into the same horizon (Gal et al., 2014, Mickler et al., 2013).

213 As Table 2 details, for eight of the CO₂ release facilities it was crucial that the injected CO₂ was released to
 214 surface. The remaining five experiments intended that the CO₂ remained in the shallow subsurface except
 215 Vrøgum which aimed to image the CO₂ but did not aim specifically for subsurface retention or surface release.
 216 In some cases the project was modified to ensure that CO₂ was released to surface as intended.

217 3.2 Experimental design

218 The site characteristics and experimental set-up varied between each of the projects that we reviewed,
 219 although some projects mimicked or built on the experimental design of other sites; for instance the
 220 Ginninderra experimental set-up is closely based on ZERT (Feitz et al., 2014a).

221 3.2.1 Site & well geometry

222 At all sites, CO₂ was delivered to the subsurface by a borehole that was either horizontal (favoured at the
 223 shallower sites), at a 45° angle (ASGARD, CO₂ Field Lab and Vrøgum) or vertical (favoured for deeper
 224 experiments; Table 2). For experiments injecting free-phase CO₂, the gas was usually released via several
 225 perforations along the pipeline rather than a single point of injection.

226 Several of the experiments have a number of physical blocks or plots for CO₂ release; the experimental area
 227 at ASGARD was divided into three blocks of eight replicate 2.5 × 2.5 m plots (Smith et al., 2013) and at PISCO2
 228 the CO₂ was injected through a grid with 16 pinholes (Gasparini et al., 2015) via a horizontal grid arrangement
 229 of thin pipes. Similarly, ZERT's 70 m long horizontal well was divided into six zones by inflatable packers
 230 (Spangler et al., 2010). The set-up at Ginninderra is similar; the 100 m long pipe is partitioned into five 16-m
 231 long segments (Feitz et al., 2014a).

232 Sites were often designed to facilitate specific needs. For example, the ZERT pipeline was laid at 45° North to
 233 maximise data resolution, and a reference grid was laid over the ground surface (Spangler et al., 2010) which
 234 subsequent projects adopted (e.g. CO₂ Field Lab, Jones et al., (2014a)). A line of perforations in the inclined
 235 well at Vrøgum aimed to simulate gas bubbling from a short fissure into flowing groundwater (Cahill et al.,
 236 2014). Finally, QICS released CO₂ to seabed via a deviated well drilled from the mainland to fulfil access and
 237 logistics criteria (Blackford et al., 2014).

238 3.2.2 Depth characteristics of CO₂ release experiments investigated

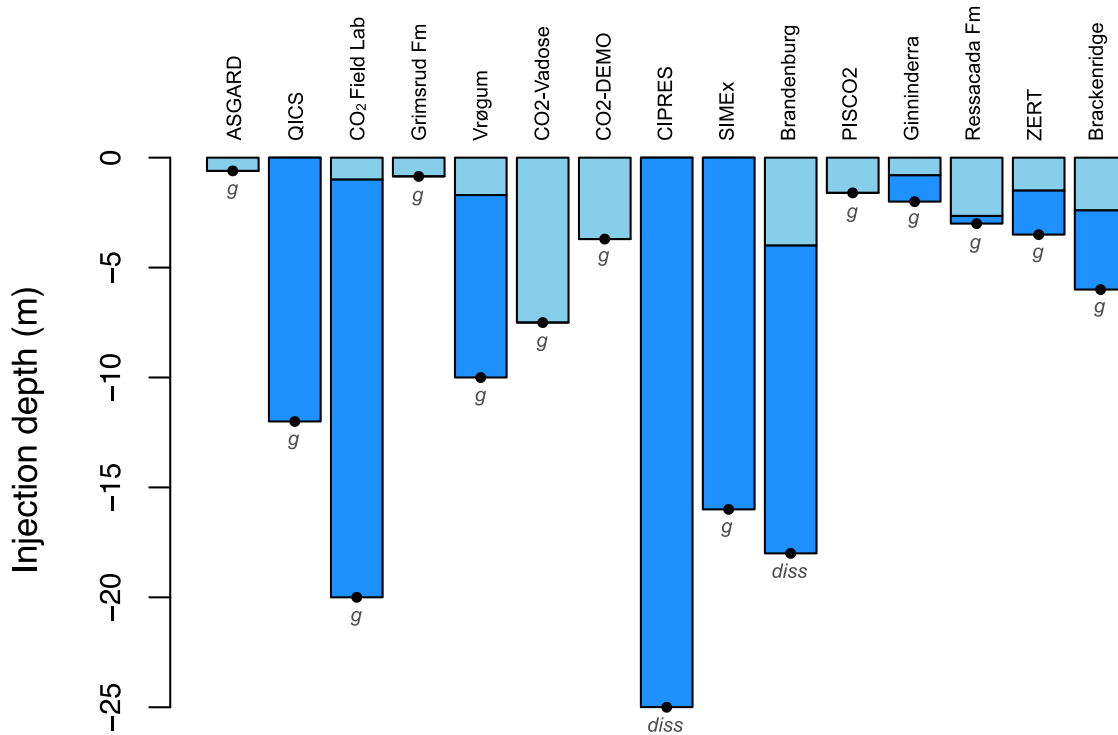
239 The CO₂ injection depths ranged from 0.6 m (ASGARD) to 25 m (CIPRES). The deepest experiment to release
 240 CO₂ to surface was CO₂ Field Lab, where the gas was injected at 20 m depth below ground level. Figure 3
 241 shows the injector depth and the maximum depth of the water table by experimental site.

242 Most projects released CO₂ into sands or gravel. CO₂-DEMO was the only shallow release experiment that
 243 injected CO₂ into a (lithified) carbonate formation. Other exceptions include ASGARD which released CO₂ into
 244 soil (Smith et al., 2013), and PISCO2 used an artificially constructed sand unit (Gasparini et al., 2015). For
 245 most projects the overburden units were the same as, or similar to, the injection formation, although thicker
 246 overburden (deeper experiments) tended to be comprised of several sediment formations making
 247 observations and interpretations of gas movement more complex.

248 The water table was relatively shallow at most of the projects, though it usually varied seasonally. For
 249 example, the water table at ZERT was less than 1.5 m, and in springtime the water table could rise to surface
 250 level (Spangler et al., 2010). Generally, the vadose zone thickness was < 4 m for all onshore experiments
 251 besides the CO₂-DEMO site, where it was ~21 m thick (Loisy et al., 2013).

252 At four sites CO₂ was always injected above the water table (ASGARD, CO₂-DEMO, PISCO2, and Grimsrud
 253 Farm), whereas at ZERT, Ginninderra and Ressacada Farm the injector could have been in the vadose or
 254 saturated zone, depending on the season. At all other sites, CO₂ was injected consistently below the water
 255 table. Most experiments (particularly those that intended to release CO₂ to surface) were conducted in the
 256 dry season, with the exception of Ginninderra, which purposefully conducted experiments in both the dry
 257 and wet season to explore the effect of seasonality (Feitz et al., 2014a).

258 The only offshore release experiment, QICS, released CO₂ into sediments 10 m below seabed.



259

260 *Fig. 3. The injector depth (black circle) and thickness of the vadose (pale blue) and the saturated zone (dark blue) for*
 261 *each CO₂ release project in our global dataset. The CO₂ injection phase is noted, where CO₂ gas (g); CO₂ dissolved in*
 262 *water (diss).*

263 3.3 CO₂ injection parameters

264 3.3.1 Properties of injected CO₂

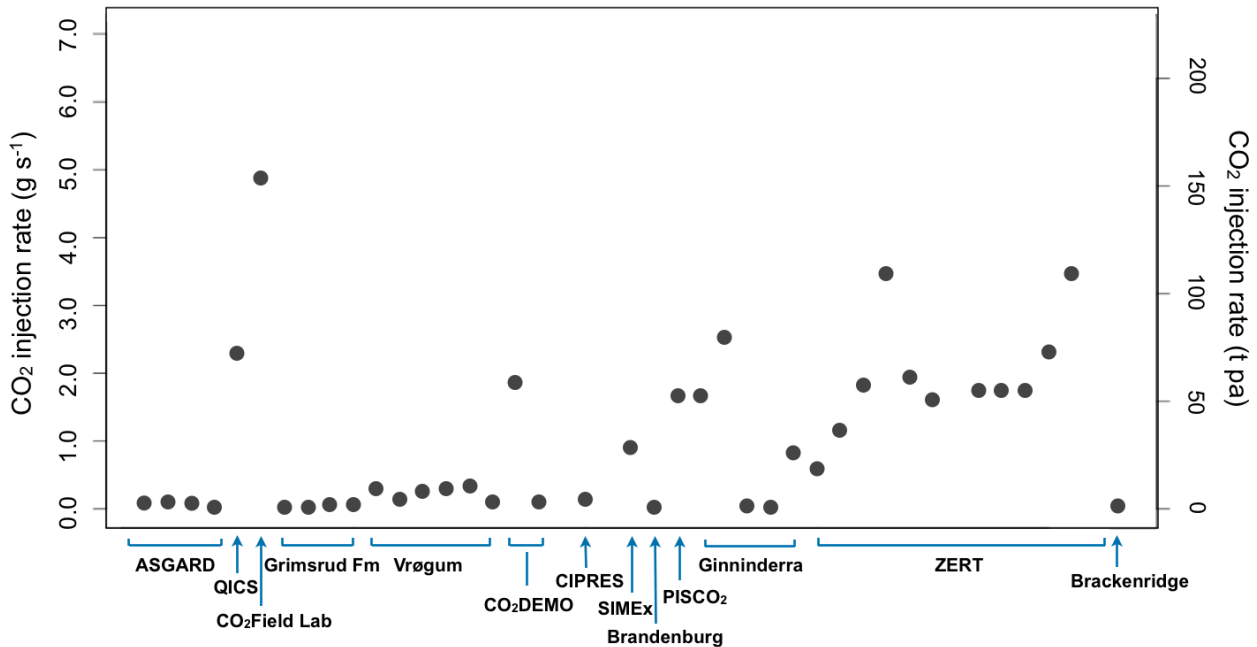
265 All experiments released free-phase CO₂ gas into the subsurface except SIMEX and Brackenridge which
 266 injected water containing dissolved CO₂.

267 All experiments used high purity (99.9%) food grade CO₂, though at CO₂-DEMO it was mixed with noble gases
 268 so that 90.57% CO₂, 5% Kr and 5% He was released (Rillard et al., 2015). Food grade CO₂ procured through
 269 chemical suppliers is often depleted in ¹³C since it is commonly sourced from processes using hydrocarbons
 270 (e.g. natural gas to urea conversion (Stalker et al., 2013)). As such, the CO₂ used for most experiments had
 271 δ¹³C values towards -30 ‰, though the exact values were site and CO₂ source specific. Carbon isotope
 272 signatures of CO₂ captured from industrial sources are typically very negative (Flude et al., 2016). Therefore
 273 the carbon isotope values for CO₂ released at field experiments can be considered representative of the CO₂
 274 that might leak from large-scale industrial stores. Possible exceptions are CO₂ that is sourced from biomass
 275 combustion, or biologically derived CO₂ which are likely to be less negative (δ¹³C(CO₂) < -6 - < -15 ‰) than CO₂
 276 derived from hydrocarbon sources (Flude et al., 2016).

277 3.3.2 CO₂ injection strategy

278 As shown in Table 2 a steady or incrementally increasing injection strategy was favoured by most
 279 experiments, though the rate varied unintentionally at SIMEX due to some challenges experienced during
 280 operation (Pezard et al., 2015b). The maximum and minimum injection rates for the CO₂ release experiments
 281 conducted at each site are shown in Figure 4. We find that there was a wide range in the rate of CO₂ injection
 282 at these experiments; the highest rate was 4.9 gs⁻¹ (153.3 t(CO₂)pa) at CO₂ Field Lab, and the smallest
 283 injection rates were 0.03 to 0.04 gs⁻¹ (0.95 and 1.3 t(CO₂)pa at PISCO2 and Brackenridge respectively). The

284 majority of experiments however inject CO₂ between 0.05 - 2 gs⁻¹, which is equivalent to 1.6 - 63 t(CO₂)pa.
 285 The injection rates were largely selected based on possible permissible leak rates from engineered storage
 286 sites and modelled properties such as injectivity (e.g. Spangler et al., (2010)). A release rate of 2 gs⁻¹ or 63
 287 t(CO₂)pa is equivalent to less than 0.001% per year of a large scale CCS project injecting 1Mt(CO₂)pa for 40
 288 years (as specified by the United States Department of Energy (US DOE) for a project aiming for over 1000
 289 years of storage), and are in the range of natural CO₂ emissions. For example, in Italy there are hundreds of
 290 natural CO₂ seeps that most commonly emit between 10-100 t(CO₂)pa (Roberts et al., 2011).

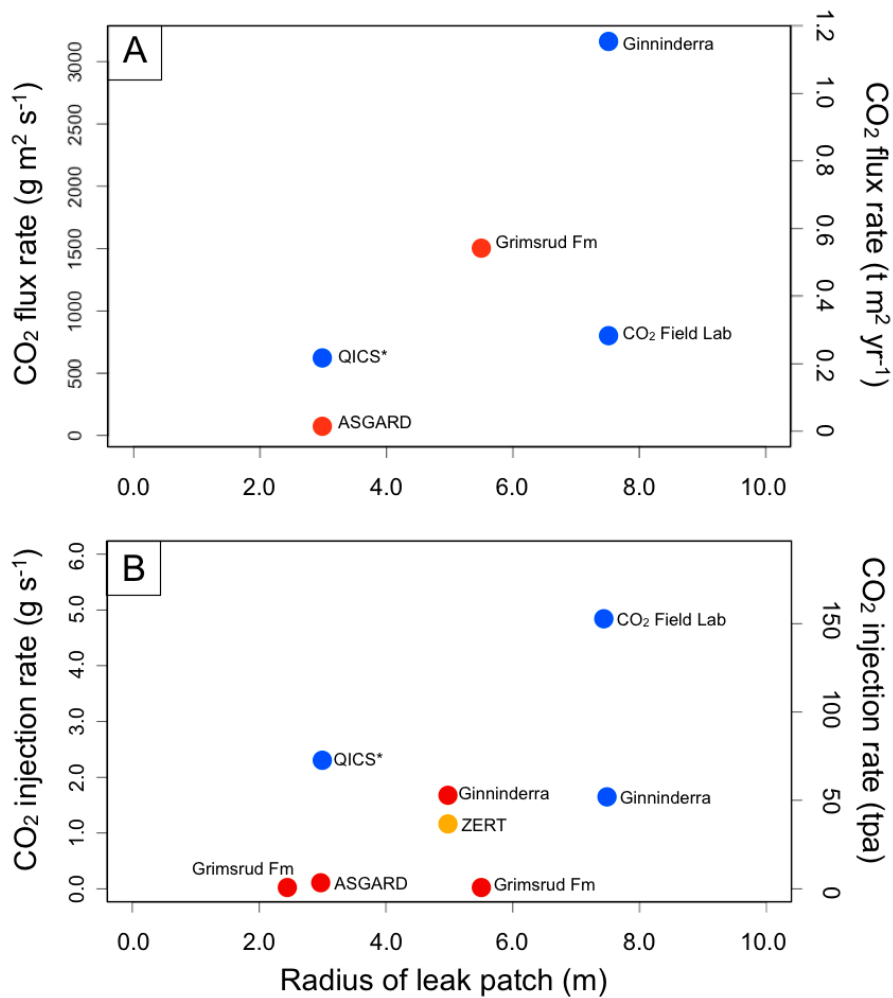


315 Where CO₂ leaked to surface, CO₂ fluxes above baseline were usually detected within 24 hours of the start
316 of injection. For example, CO₂ injected at ZERT arrived at the surface within ~5 hours (Lewicki et al., 2010),
317 and CO₂ bubble streams were observed at QICS within 3 hours of injection commencing. The greatest lag
318 time between injection and surface release was observed at Ressacada Farm and PISCO2 where surface
319 fluxes of CO₂ were not detected for 3 and 4 days respectively (Feitz et al., 2014b, Gasparini et al., 2015).
320 Unfortunately the sampling frequency was not high enough to capture the exact arrival time at a number of
321 experiments.

322 The surface leakage was typically expressed as (several) patches or ‘hotspots’ showing CO₂ flux above
323 background levels. The area of the surface hotspot was usually defined by CO₂ flux above baseline, where
324 CO₂ flux and soil gas concentrations decreased radially from maximum levels at the centre of the hotspot.
325 The hotspots that developed at CO₂ release experiments were typically between 2.5 – 5.5 m radius, though
326 leakage at Ginninderra was less patchy and occurred over a larger area than at other field sites (Feitz et al.,
327 2014a). These patches were often static once leakage had become established, for example, the leak patches
328 at CO₂ Field Lab coalesced as injection continued and then remained stable (Jones et al., 2014a). The CO₂
329 bubble streams at QICS were mobile, but concentrated in two (static) patches (Blackford et al., 2014) and the
330 characteristics of the bubble stream (e.g. bubble density and bubble size) were affected by tidally-induced
331 changes to hydrostatic pressure (Sellami et al., 2015). Patchy CO₂ leakage matches observations of leak
332 distribution at natural CO₂ seeps (Beaubien et al., 2008, Smets et al., 2010, West et al., 2015).

333 Plants and soil microbiology were found to be affected by elevated CO₂ concentrations in the soil gas and
334 land surface at the patches of CO₂ leakage. These effects were visible within a couple of days of CO₂ leakage,
335 though some plant species were found to be more resistant than others (Jones et al., 2015, West et al., 2015).
336 Observations at field and natural CO₂ release sites in a range of environments have found that soil gas
337 concentrations of 10% CO₂ and surface flux rates of 0.8 kgm²day⁻¹ is the cut-off above which the CO₂ begins
338 to impact the ecosystem (West et al., 2015). Since hyperspectral imaging can detect changes to vegetation,
339 such as leaf chlorophyll levels, these are therefore a promising remote sensing monitoring tool (Bellante et
340 al., 2013, Feitz et al., 2014b).

341 We note that there may be a correlation between the hotspot radius and the CO₂ injection rate, as shown in
342 Figure 5a, where the hotspot radius is larger at experiments with higher CO₂ injection rates. The hotspot
343 radius is not necessarily larger for greater quantities of total CO₂ injected. However we find no relationship
344 between injection rate and maximum CO₂ flux rate (graph not shown here). The flux rate and seep patch
345 radius, shown in Figure 5b, may also be correlated but there are too few data points to draw any reliable
346 conclusions. The injection depth however seems to have no effect on the CO₂ flux rate, nor does the injection
347 rate control the vertical velocity of CO₂, though we do note that the vertical flow velocity was fastest for the
348 deepest experiments where the density difference between CO₂ and surrounding pore fluids would have
349 been greatest (see Table 2).



350

351 *Fig. 5: Radius of the leakage patch and (A) maximum CO₂ flux rate and (B) maximum injection rate. The symbol colour*
 352 *indicates whether the CO₂ injection depth was into the saturated zone (blue) or the vadose zone (red) or if this was*
 353 *variable throughout the experiment (orange). The patch width appears proportional to the flux of leaked CO₂. Note some*
 354 *sites conducted more than one release experiment. QICS is starred as the leakage radius might be considered to be either*
 355 *<3 m or >5 m, since the bubbling could be interpreted to occur in two smaller patches or in one larger patch (Blackford*
 356 *et al., 2014).*

357

358 CO₂ hotspots often didn't occur vertically above the point of CO₂ injection. In some cases, the lateral
 359 migration distance of CO₂ that emerged at hotspots was greater than the depth of the CO₂ injector - even
 360 when this was shallow. For example, at Ressacada Farm the hotspot developed in a surface depression next
 361 to a road located ~30 m away from the CO₂ release point (which was at 3 m depth) (Feitz et al., 2014b). At
 362 the subseabed QICS injection site, patches of CO₂ bubble streams established ~ 10 m west of the CO₂ diffuser
 363 (Blackford et al., 2014), along strike of the sediment structure (Cevatoglu et al., 2015).

364 Despite attempts to minimise subsurface disturbance during site construction, a number of projects found
 365 the soil structure was disturbed to an extent that the characteristics or location of the resulting CO₂ leak were
 366 affected. For example, at ZERT, despite the use of directional drilling to install the CO₂ pipeline (Spangler et
 367 al., 2010), the CO₂ leakage locations corresponded to small elevations in the horizontal well. It is thought that
 368 gas collected at these high points within the pipe before leaking to the surface (Strazisar et al., 2009). CO₂
 369 released at Grimsrud Farm preferentially leaked along the border of the two artificial plots (i.e. where the
 370 soils in the experimental plot contacted with undisturbed surrounding soils) (Moni and Rasse, 2014), and
 371 preferential flow pathways developed at the artificial PISCO2 site where injection and pumping tests were
 372 performed prior to CO₂ injection (Gasparini et al., 2015).

373 3.4.2 The subsurface characteristics of CO₂ leakage

374 CO₂ was often detected in soil gas prior to surface flux, and so was usually detected within a matter of hours
 375 following the start of CO₂ injection. Towards the centre of the hotspots, soil gas could reach up to 100% CO₂,
 376 but the extent of lateral spread of CO₂ in the subsurface tended to be much greater than the footprint of
 377 surface degassing. Soil gas saturation was spatially and temporally variable, depending on weather (see next
 378 section), subsurface structure, and CO₂ flow rate. For example, at ASGARD CO₂ moved preferentially through
 379 the higher permeability sandy and gravely deposits lying beneath the injection point, rather than migrating
 380 through the overlying soils (West et al., 2009). Soil gas monitoring at CO₂ Field Lab detected CO₂ in soil gas
 381 directly above the injector shortly after injection commenced. However, there was no CO₂ flux to surface in
 382 this region; leakage was established northeast (up dip) of the injector a day later, following an increase in the
 383 injection rate (Barrio et al., 2013, Jones et al., 2014a). It is believed that CO₂ initially leaked up the well casing
 384 but then, at higher injection rates, the CO₂ favoured an alternative pathway to establish leakage away from
 385 the wellbore. Observations at QICS and CO₂ Field Lab also found that the subsurface CO₂ flow path was
 386 affected by CO₂ injection rate (Blackford et al., 2014, Cevatoglu et al., 2015). Sediment grain size controlled
 387 CO₂ flow initially at QICS until the gas pressure or gas volumes overrode the stratigraphic controls, spatially
 388 focussing the CO₂ flow via the formation of chimney structures (Cevatoglu et al., 2015).

389 In the saturated zone, maximum gas saturation in the sediments was ~7% at Vrøgum and CIPRES (Pezard et
 390 al., 2015a). CO₂ saturation at Vrøgum was found to be proportional to the grain size properties of the
 391 sediments, and so CO₂ spread was governed by the permeability and structure of the subsurface (Lassen et
 392 al., 2015). Gas concentrations were greater in higher permeability sediments, and the geochemical effects
 393 (pH, EC) of CO₂ faster and more uniform (Cahill et al., 2014, Schulz et al., 2012, Yang et al., 2015). As a result
 394 the plume spread towards regions with higher permeability, even overcoming groundwater flow to do so
 395 (Yang et al., 2015). CO₂ migration was successfully imaged by resistivity changes which are incurred as
 396 migrating CO₂ partially displaces the pore waters. For example, at Ressacada Farm resistivity changes were
 397 consistent with CO₂ leakage pathway to the hotspot location (Oliva et al., 2014).

398 Similar to the subsurface spread of CO₂ gas, the spatial extent of geochemical impacts was also much wider
 399 than the extent of surface release. For instance, bubble vents at QICS were located within 10 meters of the
 400 subseabed injector, whereas the spatial extent of the geochemical impact of the injected CO₂ in the
 401 sediments and pore waters was contained to 25m of the injection point (Lichtschlag et al., 2015). Monitoring
 402 techniques deployed at Vrøgum and Brandenburg observed a two-phase geochemical evolution of the CO₂
 403 leak, where a pulse in ion concentrations is followed by persistent acidification (Cahill et al., 2014, Yang et
 404 al., 2015). In the early stages of injection there tended to be a delay before any chemical changes were
 405 detected, thought to result from initial CO₂ gas flow in discrete channels, limiting the contact with the water-
 406 phase and so restricting the quantity of CO₂ that dissolves into the groundwater (Lassen et al., 2015). The
 407 Vrøgum researchers also noted that the unconfined aquifers were susceptible to recharge which caused
 408 rapid and inconsistent changes to the groundwater properties quite distinct from the changes caused by the
 409 presence of CO₂ (Cahill et al., 2014).

410 3.4.3 Temporal effects on CO₂ fate

411 The location, width and intensity of the hotspots that developed at onshore CO₂ release experiments were
 412 dependent on climatic conditions, including diurnal temperature (which affected wind speed), rainfall and
 413 pressure. For instance, CO₂ flux was significantly impeded by rainfall at ASGARD and temporarily stopped
 414 when the soil froze (Smith et al., 2013, Jones et al., 2014a). Generally, CO₂ flux was consistently highest in
 415 periods of low rainfall, and CO₂ soil gas concentrations were lowest in periods of low atmospheric pressure.
 416 These effects are presumably due to a combination of factors, such as the swelling and saturation of soil in
 417 wet periods, rainwater preferentially percolating via the same high-permeability pathways favoured by
 418 leaking CO₂, and the elevated pressure gradient between surface and subsurface in periods of low pressure.
 419 At the offshore QICS experiment, rates of CO₂ bubbling were also observed to vary significantly with tidally
 420 induced changes in hydrostatic pressure (Blackford et al., 2014, Sellami et al., 2015).

421 There is evidence also that longer-term seasonal variations influence the location, size and style of subsurface
 422 migration and leakage of CO₂ at artificial release experiments. This matches the seasonal changes that have
 423 been observed at natural CO₂ seeps (Heinicke et al. (2006) and references therein). For example, release

424 experiments at Ginninderra were performed in both the wet and dry seasons to explore the influence of
 425 seasons on the quantity and style of CO₂ release to surface. The injector depth at the site was such that in
 426 the wet season CO₂ was released into the saturated zone, and in the dry season, it was released into the
 427 vadose zone. In the wet season, CO₂ leaked to surface via one large patch (16 m x 30 m). This was located in
 428 a sandier region of the site, where soil gas saturation reached 80%. In contrast, leakage was more distributed
 429 in the dry season, occurring in three smaller patches which had lower and more steady-state fluxes than the
 430 single large patch that established in the wet season experiments (Feitz et al., 2014a). These hotspots
 431 established in the more clay-rich regions, and soil gas saturation in the hotspots were lower, reaching a
 432 maximum of 60%. In addition, soil surveys aided by krypton tracer found that CO₂ spread much further from
 433 the horizontal well in the wet season (Schroder et al., 2017). It is hypothesised that the greater extent of the
 434 vadose zone in the dry season allows greater quantities of CO₂ to accumulate in the subsurface, and so limits
 435 the intensity of CO₂ release to surface (Feitz et al., 2014a, Schroder et al., 2017).

436 3.4.4 Quantifying CO₂ leaked to surface

437 Quantifying the proportion of injected CO₂ that is released to surface (atmosphere or seabed) has proven
 438 very challenging at release experiments (Feitz et al., 2014b). Out of the 30 release experiments in our dataset
 439 (at nine projects) which released CO₂ to surface, estimates of quantities leaked to surface are provided for
 440 only ten experiments (at seven projects). These estimates were usually derived from measurements of CO₂
 441 flux (see Table 2) and not all reported estimates account for baseline CO₂ flux.

442 Estimates of total leakage range from 5% of the injected CO₂ at CO₂ Field Lab (reported to be a likely
 443 underestimate (Barrio et al., 2013)), up to 82-83% (at Grimsrud Farm (Moni and Rasse, 2013) and PISCO₂
 444 (Gasparini et al., 2015)) and 90.3% (for the preliminary vertical experiment at ZERT). The project researchers
 445 report that these calculations proved difficult for several reasons, including:

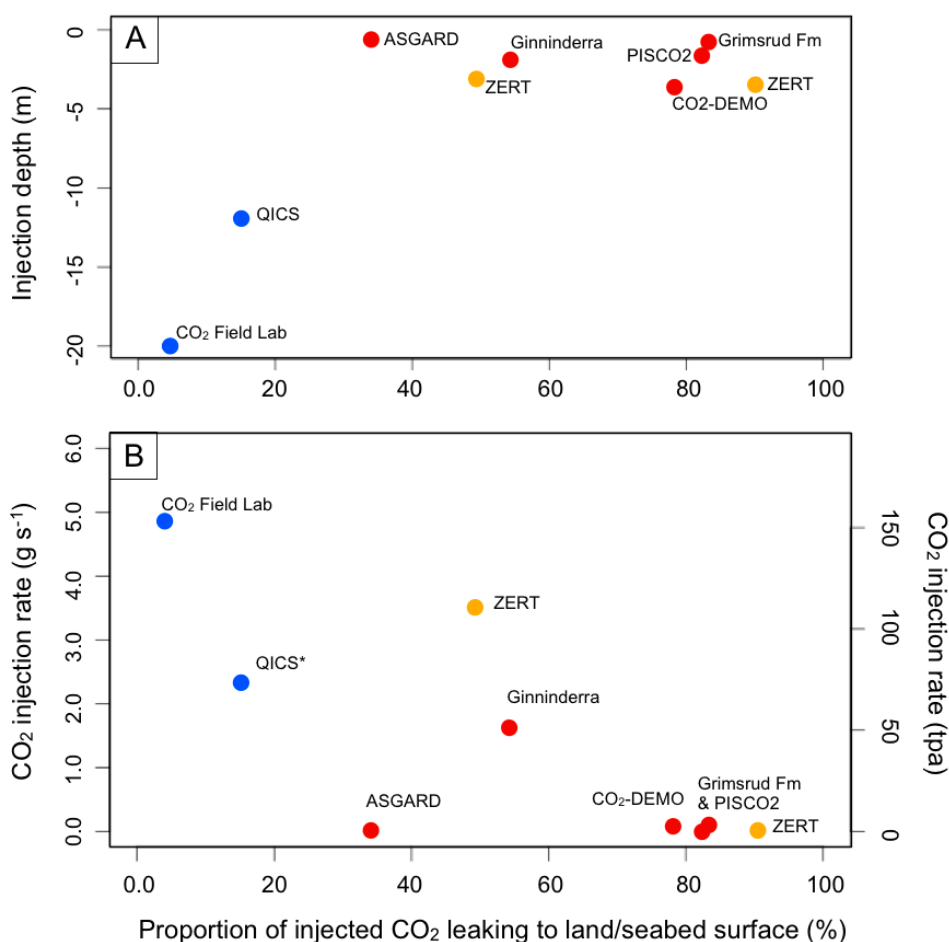
- 446 • Given the spatial and temporal distribution of CO₂ degassing, it is difficult to take continuous or
 447 detailed measurements of the seep site (Jones et al., 2014a).
- 448 • At ASGARD, ZERT and Grimsrud Farm the CO₂ migrated beyond the initial monitoring boundaries
 449 making it difficult to estimate the relative proportions of CO₂ that leaked to surface, remained in soil
 450 gas or dissolved (West et al., 2009, Lewicki et al., 2010, Moni and Rasse, 2014).
- 451 • The background soil respiration rate can be very variable.

452 These issues have led to the calculation of negative fluxes (Lewicki et al., 2010).

453 The proportion of CO₂ that leaked to surface varied in response to various environmental factors (see Section
 454 3.3.1). However it was not possible to quantitate these effects, except at QICS, where the quantity of CO₂
 455 that reached seabed as a free phase ranged from 8-15%, depending on the tide (Blackford et al., 2014, Mori
 456 et al., 2015). As such, none of the experiments that we reviewed reported how the total proportion of CO₂
 457 that leaked to surface varied as the experiment progressed. However, several experiments reported how the
 458 flux evolved as CO₂ injection continued. Usually leaked CO₂ was detected within hours after CO₂ injection
 459 started (see section 3.3.1), and CO₂ fluxes showed a steady increase until it plateaued (i.e. reached steady-
 460 state).

461 Figure 6a shows the estimated proportion of CO₂ that leaked to surface and the depth of the CO₂ injector.
 462 The two deepest experiments in the dataset (CO₂ Field Lab and QICS) leaked the smallest proportion of CO₂
 463 to surface/seabed, though as previously mentioned the value for CO₂ Field Lab is likely an underestimate
 464 (Barrio et al., 2013).

465 There is a more obvious trend between the proportion of CO₂ that is released to surface and the maximum
 466 injection rate (Figure 6b), but again, there are too few data points for this trend to be statistically significant.



467

468 *Fig. 6: The estimated proportion of injected CO₂ that leaked to land surface or seabed as a gas at the field experiments*
 469 *plotted against (A) injection depth and (B) injection rate. The symbol colour indicates whether the CO₂ injection depth*
 470 *was into the saturated zone (blue) or the vadose zone (red) or if this was variable throughout the experiment (orange).*
 471 **offshore.*

472 Several experiments have tested added tracers as a method of detecting CO₂ leakage, including
 473 perfluorocarbon (PFC) (ZERT), SF₆ (Brandenburg), noble gases (krypton, Ginninderra; helium and argon, CO₂-
 474 DEMO). For tracers to quantify CO₂ leakage they must behave predictably and preferably conservatively (i.e.
 475 mimic the CO₂ behaviour). Co-released noble gases at CO₂-DEMO included helium and argon, which both
 476 arrived ahead of the CO₂, and so behaved as precursor tracers for the leakage of CO₂ in the vadose zone
 477 (Rillard et al., 2015). Krypton co-injected at the Ginninderra tests, and PFC tracers used at the 2007 ZERT
 478 experiments were found to correlate with soil–gas results and so track the injected CO₂ (Strazisar et al., 2009).
 479 However, observations at ZERT also indicated that the distribution of PFC was also affected by small changes
 480 in topography and possibly also gas density and/or soil properties (Strazisar et al., 2009, Spangler et al., 2010),
 481 apparently differently to the CO₂. No CO₂ release experiment has yet attempted to quantify CO₂ fate or
 482 surface flux using chemical tracers.

483 3.4.5 Quantifying the fate of injected CO₂

484 Only one experiment, QICS, reported comprehensive estimates of CO₂ fate in each of the different phases of
 485 the studied environment, each quantified using different approaches. A maximum of 15% of injected CO₂
 486 was estimated to reach seabed as bubble streams, and geochemical modelling based on pore-water
 487 observations estimated that 14–63% remained dissolved within the sediment pore water (Taylor et al., 2015a).
 488 The proportion that remained trapped in the sediment (Cevatoglu et al., 2015) was imaged by repeat seismic
 489 surveys. A different remote method imaged the proportion of CO₂ trapped as free phase gas at Vrøgum.
 490 Cross-borehole ground penetrating radar used in the July 2012 experiment found 30–40% of the injected CO₂

491 volume was trapped in this way, and it was inferred that the remaining CO₂ must have dissolved or migrated
492 (Lassen et al., 2015).

493 Although there have been detailed geochemical surveys at ZERT since the 2008 field campaign, it has been
494 difficult to estimate the quantity of CO₂ that dissolved into the groundwater. Fluctuations in the water table
495 mean that CO₂ interacts with a changing column height of groundwater, and the groundwater speed at the
496 site was unknown. Further, the very action of bubbling CO₂ through the groundwater is thought to perturb
497 the groundwater flow at the site (J. Thordsen, Pers. Comm).

498 *3.5 Additional monitoring activities*

499 The design of the monitoring array tended to be site specific, but there were usually a number of monitoring
500 boreholes of various depths to obtain regular soil gas and water samples, as well as devices for measuring
501 CO₂ flux. Atmospheric monitoring methods such as Eddy Covariance towers, or geophysical tools such as
502 Ground Penetrating Radar (GPR) were used at some field release sites and for natural analogues. At the
503 offshore QICS experiment a whole suite of monitoring approaches were deployed at and below seabed, at
504 the sea surface and in the water column, using equipment installed by SCUBA divers and deployed on towed
505 platforms (Blackford et al., 2014).

506 *3.5.1 Pre-release monitoring*

507 Pre-release monitoring must be conducted to characterise the environmental baseline at the experimental
508 field site to be able to establish which changes result from CO₂ release, and which are simply due to
509 environmental variability unrelated to the release experiment. The experiences at field experiments have
510 highlighted not only the importance of rigorous baseline monitoring, but also the type of baseline data
511 collected and the appropriate time period over which baseline data should be collected. At some
512 experiments, baseline data was collected for only a couple of days prior to injection. In such cases the
513 researchers concluded that this was not long enough (Lewicki et al., 2010, Jones et al., 2014a). Subsequent
514 projects have collected more frequent or continuous baseline data for longer; for example, 2 weeks of
515 continuous/daily monitoring was conducted at QICS and Ressacada Farm. Even so, at QICS this was not
516 deemed long enough since the porewaters had not returned to baseline by the end of this period. Projects
517 such as CO₂-DEMO and Vrøggum collected baseline data at intervals over a period of ~18 months prior to CO₂
518 release to allow a longitudinal baseline to be established. The spatial extent of baseline data should also be
519 considered; for example, as Jones et al., (2014a) concludes, the location of the hotspots away from the point
520 of CO₂ release shows how the areal coverage of baseline data must be spatially adequate, so that the
521 hotspots do not establish where no baseline was collected.

522 Without adequate baseline, it becomes very difficult to, for example, report the flux of leaked CO₂. Further,
523 there is no typical or standard baseline; background CO₂ flux and its variability are unique to each field
524 experiment site. The variability of background is important to account for when interpreting CO₂ flux
525 measurements. For example, at ASGARD, background CO₂ flux could vary by 3 to 4 fold, mostly in response
526 to rainfall and air pressure (Jones et al., 2014b).

527 *3.5.2 Post-release monitoring*

528 Previous publications have seldom mentioned the value of continued monitoring of the field site once CO₂
529 injection has ceased. Post-release monitoring at field experiments provide important information about how
530 CO₂ evolves in the absence of injection pressure, which is important for understanding CO₂ dispersion and
531 fate. Further, post-release monitoring is relevant to environmental permitting at CCS projects. It is also
532 relevant to the post-closure monitoring (both pre- and post- transfer of site responsibility) (EC, 2011), that is
533 necessary to (a) verify that there is no leakage from the engineered store, or (b) to detect and quantify any
534 leak arising and to establish when the leak has stopped following remediation.

535 We find that, while most CO₂ release experiments performed some post-injection monitoring, the length of
536 monitoring period is widely variable, for example measuring for just one day only, or for only a couple of days
537 over some period after injection ended. In many cases, the post-release monitoring was not sufficiently long
538 enough to observe the decline in CO₂ concentrations to return to baseline conditions before sampling ceased.
539 In fact, the ZERT 2008 experiments, QICS, and ASGARD are the only experiments to observe (via post release

540 monitoring) the return to baseline conditions. CO₂ flux at ASGARD returned to baseline within 2-3 days (Jones
 541 et al., 2014b), whereas at ZERT it took 15 days to return to baseline at the hotspots (above the well) but only
 542 5 days to recover further (5 m) from the well (Lewicki et al., 2010). At QICS, while CO₂ bubble streams stopped
 543 shortly after the CO₂ injection stopped, concentrations of all pore water constituents returned to background
 544 values within 18 days (Lichtsschlag et al., 2015), and microbial species took 90 days to fully recover. Vegetation
 545 recovery may take longer, and be species dependent. At Vrøgum the injection lasted 72 days but
 546 groundwater was monitored for 252 days (Cahill et al., 2014), and found that most of the free phase gas
 547 dissolved into the groundwater within 20 hours following the end of CO₂ injection (Lassen et al., 2015).

548

549 **4. Discussion**

550 *4.1 Factors affecting CO₂ migration and fate*

551 The information presented in the previous section illustrates the vast contribution that shallow CO₂ release
 552 experiments have made to current scientific understanding of near surface CO₂ flow pathways and CO₂
 553 expression. We also present how observations of CO₂ leakage at field experiments (including patchy
 554 emissions, flux rates, seasonality etc.) also largely match studies of natural CO₂ seeps.

555 Preferential flow pathways channelling CO₂ leakage were found to be governed by hydrological factors
 556 (including water table depth, ground water flow, recharge, and soil properties), geomorphological factors, or
 557 recent, often anthropogenic, disturbance. Further, the migration of CO₂ and the characteristics of leakage
 558 seems to be affected by the injection rate. Since higher injection rates correlate with larger areas of leakage
 559 (i.e. greater hotspot width), higher CO₂ flux, and the proportion of CO₂ that leaks to surface, higher CO₂
 560 injection rates may encourage lateral spread of CO₂ in the subsurface, particularly where vertical migration
 561 is restricted by the soil properties.

562 Field experiments have also highlighted that CO₂ flux will vary temporally in response to environmental
 563 parameters. This has important implications for CO₂ leak monitoring, since the period of surveying of a
 564 possible CO₂ leak above a CCS site (season, before or after rainfall, in the cool morning or hot afternoon, tidal
 565 levels) may greatly influence the measured value and its eventual interpretation.

566 *4.2 Common issues at CO₂ release experiments*

567 We identified several common issues that CO₂ release projects have encountered. Some of these issues
 568 challenged the research and impacted the levels of certainty of the conclusions that could be drawn from
 569 the experimental results. Other factors simply complicated the experimental set up, and in some cases they
 570 were addressed by the project team. Importantly, none of the issues prevented useful results from being
 571 obtained from each test, and in fact often they highlighted challenges relevant to applied CO₂ leak monitoring
 572 at commercial-scale storage sites. For example, site construction and pre-release tests affected the
 573 subsurface structure of several sites, which in turn influenced the CO₂ flow pathways. While this was
 574 unintentional, this exhibits the sensitivity of CO₂ flow to the near surface soil structure. Soils become altered
 575 or compromised through various widespread practices such as farming, roads, laying of sewerage and other
 576 anthropogenic activities and also non-anthropogenic causes. Similarly, the seabed sediment structure
 577 offshore might be compositionally variable due to trawling, or due to storm disturbance or currents, or
 578 activities during site characterisation, site construction and monitoring design. As such, disturbed soil
 579 structure is not unrepresentative of the environment in which CO₂ leaks might establish.

580 Unintentional leakage along the well bore occurred at many sites (including Ginninderra, CO₂ Field Lab, ZERT,
 581 SIMEx and Brandenburg, and possibly also at ASGARD) and in many cases, corrective engineering was
 582 necessary (Barrio et al., 2013, Mickler et al., 2013, Pezard et al., 2015a, West et al., 2015). These challenges
 583 confirm previous work which identified that the well bore offers one of the most likely leak pathways at CCS
 584 projects (IPCC, 2005).

585 Knowledge exchange is extremely valuable for shaping future research, including reducing project risks. The
 586 amount of knowledge exchange between shallow CO₂ release projects is laudable. As an example, a number
 587 of researchers from the ZERT project (one of the first and longest-lived CO₂ release projects) have been
 588 heavily involved with the design and set-up of Ressacada Farm and Ginninderra. Collaborative work facilitates

589 developing solutions to problems encountered at previous experiments, and so maximises the project
590 success.

591 Table 4 summarises some of the common issues that CO₂ release experiments have encountered, and where
592 possible, recommends how future experiments might mitigate or minimise these issues.

Common issues experience at CO₂ release experiments	
Basic information	<p>To make the project more accessible for the research community, regulator and wider public, and to facilitate the comparison of research effort on this topic, it can be helpful if each project:</p> <ul style="list-style-type: none"> • Is referred to by a SINGLE designated project name or acronym. • Has a webpage or similar which hosts information about the project, such as the collective overarching aims of the research project. • Has a nominated contact for anyone wishing to enquire about the project. <p>Ideally there would be a standard unit for reporting parameters such as CO₂ flux or arrival time and so on. Currently CO₂ might be reported in a range of units, including concentration. Similarly, there is no formal definition of how the spatial extent of a CO₂ hotspot is determined; the outer limit of the hotspot is usually defined by CO₂ flux above background by an arbitrary value which may not reflect the baseline.</p>
Experiment set-up	<p>At several projects, the initial planned injection depth needed to be modified to ensure that CO₂ reached the surface. For example, injection at Ressacada Farm was planned to be deeper, but a shallower depth was chosen after a preliminary survey found that clay lenses would cause significant subsurface spread of CO₂ and may restrict CO₂ release to surface (Moreira et al., 2014a). CO₂-DEMO was a redesign of an earlier, deeper CO₂ release project where CO₂ did not reach surface. Layers of low permeability units above the injector at Vrøggum prevented CO₂ from reaching surface, and illustrated how subtle differences in lithology can significantly affect gas migration and dissolution (Cahill et al., 2014). These difficulties show the value of preliminary site investigations and tailoring the experiment to suit the subsurface properties and the experimental aims.</p> <p>At a number of sites, disturbance to the subsurface structure during site construction and pre-release tests affected the CO₂ flow pathways. While every reasonable attempt should be made to minimise disturbance, this can be difficult to avoid.</p> <p>The well bore was the primary source of unintentional leakage at CO₂ release experiments, which in some cases required corrective engineering.</p> <p>The location of seepage hotspots was hard to predict prior to CO₂ injection (Jones et al., 2014a). Sites with a number of adjacent experimental plots found CO₂ often cross-contaminated neighbouring plots (e.g. ZERT, Grimsrud Farm, ASGARD). Similarly, CO₂ breakout or injectivity often did not occur as predicted by modelling and the pre-injection knowledge of the site (Barrio et al., 2013, Spangler et al., 2010, Moni and Rasse, 2014). These complications should be considered when designing the surface monitoring array and baseline survey so that the monitoring infrastructure should cover an appropriate area, or be flexible to allow for relocation if needed.</p>
CO₂ injection strategy	<p>Where there has been only one CO₂ release experiment conducted at a site, most researchers would choose to modify their experimental design to improve the outcomes. For example, the period of injection would ideally have been longer at CO₂ Field Lab and at QICS, to allow the experiment to reach steady state (Jones et al., 2014a) and to enhance the impact of CO₂ on sediment porewaters (Taylor et al., 2015b). This shows the value of conducting multiple releases at a site, such as at Ginninderra and ZERT, since the CO₂ injection strategy can be progressively modified.</p> <p>The isotopic signature of CO₂ can vary between canisters and suppliers, depending on the process from which it is sourced. As such, where multiple cylinders/deliveries of CO₂ are used for the experiment duration, and CO₂ isotopes are used as a monitoring tool, the isotopic signature should be recorded from each CO₂ cylinder used.</p>
CO₂ fate	<p>CO₂ flux at a single measurement point can vary due to a range of environmental factors. This is important to consider when interpreting results of CO₂ flux and using measured leakage rates to</p>

estimate total leakage quantities. It also highlights the importance of developing a robust understanding of the baseline, and factors that influence the baseline contribution.

When reporting CO₂ flux, it should be clear whether the baseline has been subtracted from the measurements or not. Baseline should be subtracted before CO₂ flux values are used to estimate the proportion of leaked CO₂.

Quantifying the proportion of injected CO₂ that was released to surface (atmosphere or seabed) has proven very challenging. Efforts have also been complicated by the need to integrate measured flux with (variable) background biological CO₂ flux measurements.

Additional monitoring

At a number of CO₂ release experiments baseline surveys were not conducted for long enough to characterise the main features of the site, and post-release monitoring was not conducted for long enough to observe the return to baseline conditions.

Baseline information should include aspects such as seasonal variations in groundwater depth and speed, as well as typical background CO₂. It would be beneficial also to understand the short- and long-term effects of unusual weather patterns.

Many CO₂ release projects have not accurately established the CO₂ arrival time to surface or into soil gas. If the arrival time is an important parameter to inform the intended work at the field experiment then sampling frequency should be particularly intense during the first days of the experiment.

593 *Table 4: Common issues identified at CO₂ release experiments and recommendations for future sites to*
594 *consider during design and operation of the experiment and monitoring procedure.*

595 *4.3 Research questions for future release experiments*

596 The field experiments that we have examined here have allowed for testing of a range of monitoring
597 techniques to identify and quantify CO₂ leakage, and for factors affecting CO₂ leakage to be investigated.
598 However, knowledge gaps remain. Current sampling approaches are high intensity, and the quantification of
599 any leakage, as required by guidelines and legislation for CCS (Dixon et al., 2015), has proven difficult. These
600 experiments have therefore illustrated the need to develop more effective and low-cost detection and
601 quantification techniques, and methods that are viable on the scale of the monitoring interval of CO₂ stores
602 and over the time frame of the storage project. Release experiments present the opportunity to test
603 promising candidate cost-effective techniques such as remote monitoring approaches (both airborne and
604 ground-based (Bellante et al., 2013), acoustic methods (Bergès et al., 2015) and chemical tracers for
605 quantifying CO₂ fate and leak rate (Myers et al., 2013, Roberts et al., 2017a). The latter have not yet been
606 used to calculate CO₂ flux. As mentioned previously, the nature of the period of recovery following CO₂
607 release has been little explored in the experiments that we investigated, and this is important for obtaining
608 and complying with legislative permits.

609 There has currently been little variation in the physical setting of the CO₂ experiments conducted to date, in
610 terms of ecosystem, topography, and subsurface properties. The experiments that were available to be
611 reviewed largely investigated the effect of CO₂ on grassland, crop or microbial species. Future experiments
612 are being developed and conducted to address broader impacts, such as that at the Environmental Impact
613 Test (EIT) facility developed by K-COSEM (Lee et al., 2016). Other ecosystems with larger or perennial species,
614 natives versus introduced species could be investigated and research could explore how ecosystems might
615 influence the soil structure or properties, and so CO₂ distribution and leakage.

616 The CO₂ Field Lab and CO₂-DEMO release experiments took place in disused quarries, but all remaining
617 onshore experiments were located on relatively flat and exposed land, and so the effects of local topography
618 have been little explored. Slight topographic variation at the ZERT site influenced the distribution of PFC
619 tracers (which are denser than CO₂), and CO₂ leakage established at a surface depression at Ressacada Farm
620 (Oliva et al., 2014). Topography is observed to influence the location and characteristics of natural CO₂ seeps
621 due to corresponding changes in depth to the water table (Roberts et al., 2014), and also atmospheric
622 dispersion of CO₂. Wind speed largely controls atmospheric CO₂ dispersion, and can also affect the accuracy

623 of CO₂ measurements. Topographic depressions or enclosed areas are usually more sheltered, and so can
 624 facilitate gravity-driven ponding of CO₂ to concentrations that are dangerous to human health (Roberts et
 625 al., 2011, Smets et al., 2010). While most onshore CO₂ storage projects are/were in relatively uniform
 626 topographic settings (e.g. Quest, Boundary Dam, and In Salah are all largely located in relatively flat
 627 environments), projects might be developed in more topographically variable terrains. As such, the effect of
 628 topography and annual weather conditions on the spread of CO₂ and tracers could be explored in more detail
 629 at future release experiments.

630 The hydraulic gradients at the sites are largely representative of groundwater flow systems in unconsolidated
 631 sandy aquifers with modest rainfalls (Lee et al., 2016). Only one experiment in the dataset, CO₂-DEMO,
 632 injected CO₂ into lithified rock. Only three field experiments were excluded from our dataset and the analyses
 633 presented here, Plant Daniel and Cranfield were not included as these injected CO₂ deeper than ~40 m and
 634 did not intend to study the migration pathways of the injected CO₂. Future experiments could release CO₂ at
 635 greater depths, and into consolidated rock formations with the aim to monitor CO₂ fate and spread and for
 636 the CO₂ to leak to surface. While deeper wells are more expensive to drill, and deeper injection depths
 637 increase the risk that CO₂ will not reach the surface, the viability of these settings could be considered for
 638 future, longitudinal experiments.

639 The collective learning and common issues at these CO₂ release experiments are therefore extremely
 640 valuable for informing future research needs. Table 5 integrates the knowledge gaps discussed above
 641 together with other challenges identified in this review.

642

Variable	Suggestions for future experiments
Experiment set-up and site information	<ul style="list-style-type: none"> • Environment: More experiments need to be conducted offshore (only one offshore experiment conducted to date). A second sub-seabed CO₂ release experiment is planned as part of the H2020-funded STEMM-CCS project (reference 654462). For onshore experiments, sites should be more topographically varied and use ecosystems other than grassland. • Subsurface properties: Release into consolidated rock, or heterogeneous subsurface structure, or carbonate units.
CO₂ injection (for each experiment)	<ul style="list-style-type: none"> • Injection depth: inject CO₂ deeper to explore CO₂ migration in more consolidated or vertically heterogeneous units. • Injection period: inject CO₂ for longer periods (months). • Properties: Record $\delta^{13}\text{C}$ of all CO₂ injected (each vessel/cylinder delivered)
CO₂ Fate	<ul style="list-style-type: none"> • Aim: to quantify CO₂ leakage • Investigate how one or more of the following affects the spread and fate of CO₂ (and also use of possible tracers for CO₂): topography, rock type, water table depth (i.e. seasons).
Monitoring	<ul style="list-style-type: none"> • Baseline surveys should be detailed and collect information over a longer period than previous experiments. • Test cost-effective detection and quantification techniques, such as remote detection methods or chemical tracers for leakage. • Post-release monitoring should be detailed and collect information for longer periods so as to observe return to baseline conditions, and ecosystem recovery following CO₂ release/mitigation.

643 *Table 5: Knowledge gaps that our synthesis has identified and that future CO₂ release experiments should*
 644 *seek to address.*

645

646 5. Summary and Conclusions

647 Geological CO₂ storage sites are selected and engineered to minimize risk of leakage and maximise long-term
 648 storage potential, and are governed by legislation to this effect. Leakage to surface may not be anticipated,
 649 but effective site monitoring requires approaches that are capable of enabling any CO₂ leaks to be identified,

650 attributed, and quantified. In the absence of leakage from commercial CO₂ stores, field-scale shallow CO₂
651 release experiments provide opportunity to develop scientific understanding of the characteristics of CO₂
652 leaks that might arise from CCS projects, and methods of detecting them. We have collated a global dataset
653 of experiments conducted between 2006 - 2017, drawing on information in the published domain
654 complemented by correspondence with researchers from specific sites. The experimental procedure and
655 results were then compared to establish the collective knowledge gained from these global efforts to date.
656 We also determine whether current capabilities are fit for purpose, and if not, how future research can
657 address the remaining knowledge gaps.

658 In this way, we examined 14 different CO₂ release projects, at which a total of 42 different CO₂ release
659 experiments have been conducted. Other controlled release sites where injection depth was greater than 25
660 m were considered but not included in this paper. Collectively these experiments released 82.8 tonnes CO₂
661 over 994 days. Nine of the CO₂ release facilities intended that the injected CO₂ be released to surface, the
662 majority of remaining experiments intended for CO₂ to remain in the shallow subsurface - usually to
663 investigate groundwater interactions. The experiments show a range of test approaches, including CO₂
664 release duration, modes of release (horizontal, angled, vertical pipes), and injection depths.

665 We have identified a number of common perils and pitfalls, which future experiments should seek to learn
666 from and avoid. The main issues include leakage of CO₂ along the wellbore or pipeline, which in some
667 instances required action to remediate, and disturbing the subsurface during construction of the
668 experimental site in such a manner that these changes influence CO₂ spread and leakage. Many experiments
669 also faced challenges around determining the appropriate area of monitoring, which rarely matches the
670 models.

671 Our comprehensive synthesis has shown how the collective experience at the field experiments conducted
672 to date have advanced current knowledge of near surface flow pathways, environmental impacts, and
673 methods of detecting CO₂ leaks. The importance of establishing baseline conditions for an appropriate time
674 period cannot be underestimated, since this is important for estimating CO₂ impacts, fate, flux rates and total
675 CO₂ leakage. Crucially, quantification of any leakage has proven difficult, despite intensive monitoring using
676 multiple approaches at a number of the sites. Further work is needed for any leaks to be quantified to the
677 degree of confidence acceptable for relevant legislation. Cost-effective approaches for doing so include
678 remote sensing methods or mobile devices, or the use of chemical methods such as isotopic tracers, and
679 shallow CO₂ release experiments provide excellent opportunity to trial these methods. Tests in more
680 consolidated rock environments and carbonate hosted environments are recommended, to diversify
681 observations. Further, only one of the projects that we reviewed was located offshore (QICS), highlighting
682 the need for more activity on CO₂ leakage into the marine environment.

683

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707

708 **8. Data Availability**

709 The detailed dataset of variables outlined in Table 1 are provided via the UKCCSRC Data and Information
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Table 2. Compilation of controlled CO₂ release experiments around the world conducted to date. Half of the sites have conducted more than one release experiment ('No. of exp'), some of them preliminary tests, and the experiment phases often differed in the length ('inj. length (days)') and rate of CO₂ injection ('Max inj. rate t(CO₂)pa'). The style of the injection also varies between sites, injecting CO₂ as a gas (g) or dissolved in water (diss), and via an inclined, vertical (v) or horizontal (h) well, and at a steady, incremental (incr.) or variable (var.) injection rate.

N	Name or Acronym	Country	No. of exp	Year(s)	Inj. depth (min-max, m)	CO ₂ phase	Well orientation (degrees from vertical)	Inj. style	Inj. length (min-max, days)	Max inj. rate t(CO ₂)pa	Surface leakage? (time*)	% of CO ₂ leaked (method of derivation)
1	ASGARD	UK	4	2006 - 2010	0.6	g	45°	Steady	14-16	3.1	Y (<24 h)	34% (measured)
2	QICS	UK	1	2012	12	g	90° (h)	Incr.	36	72.53	Y (<3 h)	15% as bubbles (measured)
3	CO₂ Field Lab	Norway	1	2011	20	g	45°	Incr	5	153.3	Y (24 h)	5% (measured)
4	Grimsrud Farm	Norway	4	2012	0.85	g	90° (h)	Steady	75	1.93	Y (<12 d)	82% (measured)
5	Vrøgum	Denmark	6	2012	5-10	g	45°	Incr	2-72	4.3-10.51	N	
6	CO2DEMO	France	1***	2010-2014	3.7	g	0° (v)	Steady	<1	3.06	Y (14 h)	78% (measured)
7	CIPRES	France	2	2013	25	diss	0° (v)	Steady	2	4.38	N	
8	SIMEx	France	1	2013	13-16	g	0° (v)	Var.	0.1	550.6	N	
9	Brandenburg	Germany	1	2011	18	g	0° (v)	Steady	0		N	
10	PISCO2	Spain	1	2012	1.6	g	90° (h)	Steady	46	0.96	Y (3 d)	82.3% (modelled)
11	Ginninderra	Australia	4	2010	2	g	90° (h)	Incr / steady	56-80	21.8-79.6	Y (<24 h)	54% (measured)
12	Ressacada Farm	Brazil	1	2013	3	g	0° (v)	Incr	12	1.31	Y (10 d)	
13	ZERT**	USA	13	2007-2014	1.1-2.5	g	90° (h)	Steady / var.	7-10	0.95-110.4	Y (<24 h)	49% (measured)
14	Brackenridge	USA	2	2011-2012	6	diss	0° (v)	Steady	2		N	

* time lag, in days (d) or hours (h), between the start of CO₂ injection and detecting CO₂ leakage at the land or seabed surface.

**An experiment was conducted at ZERT in Autumn 2006 where CO₂ was released from a vertical pipe for 10 days to simulate leakage from well failure. All subsequent experiments were from a horizontal buried pipe, designed to simulate leakage via a line source such as a fault or fracture. Unless explicitly stated, in this paper, the ZERT facility refers to the horizontal injection experiments.

*** The CO₂-DEMO project followed the CO₂-VADOSE project at the same site. The deeper CO₂-VADOSE experiment did not observe CO₂ release to surface. The site was subsequently modified to ensure that surface release would occur for the CO₂-DEMO project. We refer only to the CO₂-DEMO release experiment here.